



Chrono-biostratigraphic analysis of the Lower Pleistocene paleotsunami deposits overlying the Late Pliocene Erosional Surface (LPES) at the Manavgat Subbasin (SW Türkiye) based on new nannofossil and fossil ascidian spicule data

ENİS KEMAL SAGULAR¹ Orcid: 0000-0003-0853-3912

GÜLİN YAVUZLAR-ÖZER² Orcid: 0000-0002-5042-7227

¹ Süleyman Demirel University, Faculty of Engineering, Department of Geological Engineering, 32260 Isparta, Türkiye

² Isparta University of Applied Sciences MS, PhD candidate Micropaleontology 32260 Isparta, Türkiye

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ÖZ

Doğu Akdeniz kenar havzalarında yer alan Antalya Havzası'nın doğu kesimini oluşturan Manavgat Alt Havzası'ndaki Geç Pliyosen regresyonu ve erozyonu dönemi sonrasında oluşan bir aşınma yüzeyi tespit edilmiştir. Bu hiyatüs, olasılıkla Akdeniz'in tabanında meydana gelen ve kıyı düzlüğünde hızlı ve kısa süreli deniz taşkınlarına neden olan depremlerin sonucu oluşan Alt Pleistosen paleotsunami çökelleri tarafından örtülmektedir. 10 inceleme noktasından alınan kayaç örneklerinin sedimantolojik ve biyostratigrafik incelemesi, bu yeni hızlı transgresif istifin kanıtlarını sağlamıştır. Alt Pleistosen paleotsunamit zonu bölgede, 2-3 m kalınlıkta uzanımına sahiptir. Çamurlu bir matris içinde kumdan blok boyutuna kadar değişen tanelerden oluşan kötü boylanmış kırıntılı çökellerden oluşur. Çamurlu matris, farklı ortamsal, sedimantolojik ve stratigrafik kökenleri temsil eden çeşitli makro ve mikro fosil verilerini içermektedir. Kumdan blok boyutuna kadar olan bileşenler, havza dışı sedimanter istiflerden yeniden çökeltilmiş farklı yaşta marn, kumtaşı, çakıltası ve kireçtaşı parçalarından ve havza içinden taşınan çökelimle eşyaşlı resif bloğu veya kayaç parçalarından oluşur. Tespit edilen kalkerli nannofosil verileri, Miyosen yaşlı olarak bilinen Karpuzçay Formasyonu'nda iki farklı yaşta tortul istifi ortaya çıkarmıştır: (1) Altta Erken Pliyosen yaşlı kırıntılı şelf çökelleri "Karpuzçay Formasyonu" ve (2) "Manavgat Formasyonu Tsunami Üyesi"nin paleotsunami çökelleriyle başlayan sığ denizel ve kıyısız kırıntılı çökelleri içeren, Erken Pleistosen yaşlı "Manavgat Formasyonu". Karpuzçay Formasyonunun mavi-gri marnları, NN15 - Reticulofenestra pseudumbilica zonunun (geç Erken Pliyosen) %45-60 oranında çökelimle eşyaşlı nannofosil topluluğuna sahiptir, Manavgat Formasyonunun gri-

kahverengi çamurtaşları ise NN19a - *Gephyrocapsa caribbeanica* Astzonunun (Erken Pleistosen) %5-15 oranında çökelimle eşyaşlı nannofosil verilerini içerir. Manavgat Formasyonu Tsunami Üyesi tortullarına ait nannofosil verileri, tsunami olayının 1,81 My ile 1,60 My (Gelasian-Calabrian geçişi) arasında meydana geldiğini göstermektedir.

Anahtar Kelimeler: Ascidian spikülü, Erken Pleistosen, Manavgat Alt-havzası, nannofosil, tsunami tortulları

Enis Kemal Sagular eniskemal@gmail.com

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ABSTRACT

*In the Manavgat Subbasin, which is the eastern part of the Antalya Basin located at the Eastern Mediterranean marginal basins, an erosional surface that formed after the late Pliocene regression and erosion period was determined. This hiatus is overlain by Lower Pleistocene paleotsunami deposits that are probably the result of earthquakes below the Mediterranean that caused rapid and short-lived sea flooding of the coastal plain. Sedimentologic and biostratigraphic examination of rock samples, taken from 10 locations, provided evidences of this new regressive succession. The lower Pleistocene paleotsunami zone extends in 2-3 m. thickness at the region. It consists of poorly sorted clastic sediments which are composed of sand to boulder-sized grains within a muddy matrix. The muddy matrix contains various macro- and microfossils data representing different environmental, sedimentological and stratigraphic origins. The sand to boulder components are composed of different aged marl, sandstone, conglomerate, and limestone fragments, which were reworked from extrabasinal sedimentary successions and coeval reef block or boulders that were removed from the intrabasin. Detected calcareous nannofossil data have brought out two different aged sedimentary sequences in the Karpuzçay Formation known as Miocene aged: (1) Early Pliocene aged clastic shelf deposits "Karpuzçay Formation" below and (2) Early Pleistocene aged "Manavgat Formation" contains shallow marine and coastal clastic deposits that begun with the paleotsunami deposits of "Manavgat Formation Tsunami Member". The blue-grey marls of the Karpuzçay Formation have 45-60% syndimentary nannofossil assemblage of NN15 - *Reticulofenestra pseudoumbilica* zone (late Early Pliocene), grey-brown mudstones of the Manavgat Formation include 5-15% syndimentary nannofossil data NN19a - *Gephyrocapsa caribbeanica* subzone (Early Pleistocene). The nannofossil data of Manavgat Formation Tsunami Member deposits suggest that the tsunami event occurred between 1.81 Ma and 1,60Ma (Gelasian-Calabrian transition).*

Keywords: Ascidian spicule, Early Pleistocene, Manavgat subbasin, nannofossil, tsunami deposits

1. INTRODUCTION

In this study, the presence of an unexplored angular unconformity within the Karpuzçay Formation which was defined by Akay et al. (1985) in the Manavgat Subbasin (Antalya) of south-western Turkey is discussed. Although previously described as Aksuçay Formation (Akbulut, 1977), Manavgat break (Monod, 1977), Beşkonak Formation (Eroskay, 1968) after then marine sedimentary rock units comprising shale-sandstone-conglomerate alternation with rare volcanic tuff interbeds were redefined as Karpuzçay Formation limited biostratigraphic data by Akay et al (1985). According to these previous researchers, this sedimentation period was assumed between Upper Oligocene and Upper Tortonian. Most researchers such as Monod, O., 1977; Poisson, 1977; Akay et al.,1985); Şenel (1997), Flecker et al. (1998), Robertson (1998), Karabıyıköğlü et al. (2000), Robertson et al. (2003), Flecker et al. (2005), Karabıyıköğlü et al. (2005), Çiner et al. (2008) reiterated that Karpuzçay Formation represents to be a continuous, regressive Miocene deposition. First, Sagular (2009) indicated that marine clastic sedimentary rocks, including tsunami deposits of the Karpuzçay Formation, represent the Plio-Quaternary sedimentation based on the nannofossil and didemnid records in the study of the Manavgat region. However, the author neither separates the different sedimentary parts of these sediments into the Pliocene and Pleistocene sequences, nor did also perform a detailed biostratigraphic study. The necessity of dividing Karpuzçay formation into two different formations (Karpuzçay Formation-1 and Karpuzçay Formation-2) was firstly proposed by Sagular et al. (2016). In Sagular and Yavuzlar (2017), tsunami deposits within the Karpuzçay Formation as Karpuzçay Fm-2 Tsunamite Member in detail. The authors revealed an erosional surface that formed after the late Pliocene regression in the Manavgat Subbasin. They described that this unconformable surface was covered by lower Pleistocene paleotsunami deposits probably occurred by rapid and short-lived sea overflows on the coastal plains because of earthquakes at the base of the Mediterranean. According to Herman et al. (2015), these earthquakes have continued as significant rates of moderate level seismicity at the southern edge of the Anatolian Block laid the east-west trending Cyprian Arc, representing the convergent boundary between the Anatolian and the African Plate for 50 million years. Sedimentary formations similar to those described in this study, where mud and megablocks are deposited together, where deep and shallow marine sediments and fossils are present coexistent, are described in detail as earthquake tsunami deposits by different researchers such as Dawson and Stewart (2007), Fujiwara and Kamataki (2007), Kortekaas

and Dawson (2007), Morton et al. (2007). In this study, we have been dealt with the nannofossil contents of the shallow marine sediments which occurred with a new rapid transgression initiated by an earthquake and tsunami event at the Eastern Mediterranean in the early Pleistocene.

2. MATERIAL AND METHODS

In the study, comparative biostratigraphic proofs were obtained from many different micro- or macrofossil data such as nannofossils, ascidian spicules, planktic and benthic foraminifers, molluscs, reefal fragments or the others (i.e. honey bee remnant) in the 47 rock samples collected from 10 investigation spots. Biostratigraphic data of tsunami levels were obtained from nannofossils and ascidian spicules in smear slides conventionally prepared to Perch-Nielsen (1985a), Bown and Young (1998) and thin sections prepared to Sagular (2003) method were examined under the polarised light microscopes. By means of the tip of a sterile needle is made scrapping from the fresh fracture surface of the sample onto a microscope slide. Scraped sample solution made with one or two pure water drops is dried on a hotplate. The slide is covered with a coverslip having a drop of Canada balsam heated on the hotplate. Then prepared smear slide is examined by means of Nikon and Leica polarizing research microscopes, under 1000x or more magnifications. In thin sections prepared from carbonate-rich sediments may be enabled to search of nannofossils but need to be used thin section preparation methods as explained by Bown and Young (1998), Sagular (2003). Thin sections allow making a distinction between synsedimentary and reworked fossil records in thin-grained limestone or arenites. Applied imaging techniques under five transmitted light types such as normal light (NL), polarized light (PL), phase contrast (CL), polarized light with gypsum plate (GL), and with quartz plate (QL) were taken from Martini (1971), Perch-Nielsen (1985a), Bown and Young (1998), Sagular (2003). The semiquantitative abundance of synsedimentary or reworked calcareous nannofossils was estimated as the number and proportion of species for 200 field of views in the smear slides in accordance with Bown and Young (1998), and Sagular (2003). However, in the fine-grained calcareous clasts and calcareous matrix/cement of the thin sections were merely sought out the existence of the mainly reworked nannofossil species in accordance with Sagular (2003). Nannofossils species and biozones referred to Martini (1971), Okada and Bukry (1980), Perch-Nielsen (1985a, 1985b), Young et al (2017). In

addition, for the identification and dating of the ascidian spicules were utilized by Varol and Houghton (1996) and Sagular et al (2018).

3. GEOLOGICAL SETTINGS

In the study area 5 geologic units were defined: (1) Oymapınar Limestone which is consist of yellow-cream colored reefal or sparritic limestones was named after Oymapınar village on the edge of Manavgat River (Monod, 1977), (2) Geceleme Formation which comprises cream-brown marls interbedded/alternated calcarenite/limestone (Akay et al, 1985), (3) Karpuzçay Formation which is composed of blue-green-grey marl and limestone or marl-claystone-sandstone-conglomerate alternations (lower part of the formation identified by Akay et al, 1985), (4) Manavgat Formation which includes grey-brown mudstone-sandstone alternations and upper deltaic coarse clastics (upper part of Karpuzçay Formation and completely Taşlık Formation identified by Akay et al., 1985; new suggestion in this study), (5) Upper Quaternary deltaic, coastal and terrestrial deposits (Fig. 1). Oymapınar Limestone, Geceleme and Karpuzçay Formations referred in the study were dealt for the Aksu, Köprüçay and Manavgat Basins by Akay et al. (1985), Akay and Uysal (1985). Except Sagular (2009), ages of these same kind marine sediments, including also tsunami deposits in the study area, in the Aksu, Köprüçay, and Manavgat basins, were reported as Miocene in former studies (Gutnic et al., 1979; Akay et al., 1985; Flecker et al., 1998; Robertson, 1998; Karabıyıkoglu et al., 2000; Robertson et al., 2003; Flecker et al., 2005; Karabıyıkoglu et al., 2005; Çiner et al, 2008). Sagular (2009), in the study of the Manavgat region, based on the nannofossils and didemnid records, determined that marine clastic sedimentary rocks involving tsunami deposits of the Karpuzçay Formation represent the Plio-Quaternary sedimentation. However, the author did not separate the different sedimentary parts of these sediments into the Pliocene and Pleistocene sequences, nor did also perform a detailed biostratigraphic study.

Akay et al. (1985) defined the marine sedimentary rock units theoretically or based on less and relative biostratigraphic data as Neogene-Quaternary sedimentary successions. Authors described that the Antalya Miocene basin in the west of Middle Taurus lies between the Beydağları and Anamas-Akseki platforms; the basin consists of the Aksu Formation which includes terrigenous conglomerate-siltstone, marine conglomerate-sandstone, and reef limestone lenses; the Oymapınar Limestone of mainly reef limestone;

the Çakallar Formation consisting of limestone breccia and packstone alternated with clayey limestone; the Geceleme Formation of limy claystone-sandstone alternation; and the Karpuzçay Formation which is composed of shale-sandstone-conglomerate alternation with occasional volcanic tuff interbeds.

According to the authors, this sedimentation period was from Upper Oligocene to Upper Tortonian. And they introduced that the Taşlık Formation consisting of clayey limestone-limestone-blocky conglomerate (some are gypsum) deposited locally in Lower Messinian. Although they could not accurately prove nannofossil zones between NN7 and NN11, in their study, they argued that all standard Miocene nannofossil biozones from NN1 to NN11 exist in the Antalya Basin. Afterwards many researchers such as Flecker et al. (1998), Glover and Robertson (1998a,1998b), Robertson (1998), Karabıyıköğlü et al. (2000), Poisson et al. (2003), Robertson et al. (2003), Deynoux et al. (2005), Flecker et al. (2005), Karabıyıköğlü et al. (2005), Çiner et al. (2008) also asserted that are present the same stratigraphic structure in their numerous regional geologic, tectonics, sequence stratigraphic or sedimentologic studies in the Manavgat Basin, based on the stratigraphic findings and interpretations of Akay et al. (1985).

Tsunami deposits were identified as F11: Massive pebbly mudstone by Karabıyıköğlü et al. (2005) or were described as F10: massive pebbly mudstone facies by Çiner et al. (2008) within the Karpuzçay Formation. However, these sedimentary levels belonging to the Karpuzçay Fm were accepted as tsunamite in the studies of Flecker et al. (2005), Sagular (2009).

As referred to in this study, in the other study of the authors (Sagular and Yavuzlar, 2017) tsunami deposits have been proposed to be named as Manavgat Formation Tsunamite Member. In addition, the upper part of the (formerly-named by Akay et. al, 1985) Karpuzçay Formation, which includes tsunamite level in its below, has been renamed as Manavgat Formation.

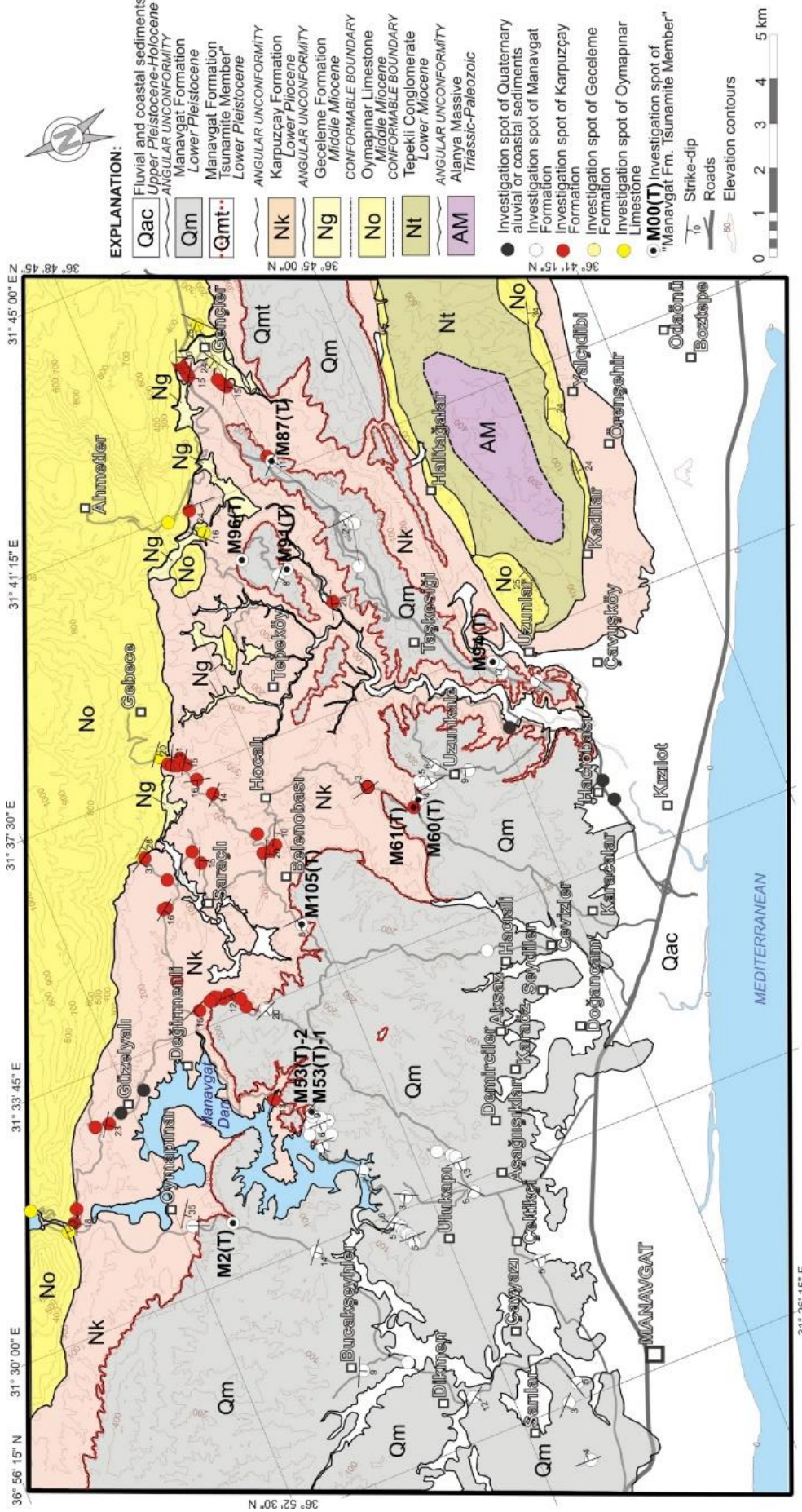


Figure 1. Geological map of the Neogene–Quaternary sedimentary successions prepared in the study area located at the Antalya–Manavgat Subbasin (SW Turkey).

Şekil 1. Antalya–Manavgat Alt Havzasında (GB Türkiye) yer alan çalışma alanında hazırlanan Neojen–Kuvaterner tortullarına ait jeolojik harita.

Four stratigraphic sequences are placed in the Manavgat Subbasin (Fig. 2): (1) At the bottom, two conformable successions, comprising Oymapınar Formation (Langhian) and Geceleme Formation (Langhian-Serravalian), involve in the first sequence. Oymapınar formation comprises yellow-cream colored, fossiliferous, thick-stratified, sparritic limestones deposited in carbonate shelf or slope of neritic environment. Geceleme Formation consists of alternations of cream-light brown limestone/calcarenite and light brown marls occurred in outhter (carbonate) shelf. (2) The second sequence is the Karpuzçay Formation including Early Pliocene aged clastic and carbonate facies sediments. At the bottom, overlaying the Geceleme Formation with angular unconformity, Karpuzçay Formation begins the alternations of red-green limestone-calcarenites and blue-green marls, representing clastic to carbonate shelf environments and Pliocene transgression in the northeast coasts of the Eastern Mediterranean. Upwards, it presents blue-green siltstone/marls alternated with red-brown sandstones belonging to offshore transition and clastic shelf and a regressive character. In the regressive upper levels, green-grey mudstones alternated with common burrows-having, wavy/lenticular cross-bedded, red-brown sandstones represent a transition between offshore shelf and lower shore-face. In the top levels of the sequence, the formation consists of highly burrowed, wavy/flaser cross-bedded, green-grey mudstones with coal lenses alternated with red-brown sandstones regressively deposited in upper delta / tidal and coastal flat in the subbasin. It is observed that the sequence which had begun with transgressive character then it ended in regressive.

The maximum thickness of the unit was measured approximately 750 meters in the study area. (3) Third sequence begins with the tsunami deposits (Manavgat Formation Tsunamite Member) taken in place bottom levels of the Manavgat Formation and overlays angular-unconformably the Late Pliocene Erosional Surface in the Manavgat Subbasin. Tsunamite zone includes multigenetic pebbly mud matrix involving recrystallized limestone boulders/blocks/cobbles (Mesozoic aged), conglomerate-sandstone blocks/cobbles (Neogene aged), marl blocks (Early Miocene or Early Pliocene aged), reef boulders/cobbles and muddy rip-up clasts (cobble-pebble sized). Sedimentary formations similar to those defined in this study, where mud and megablocks, deep and shallow marine fossils and clastics are deposited together, being poorly sorted, irregular successions, were described in detail as earthquake tsunami deposits by Dawson and Stewart (2007), Fujiwara and Kamataki (2007), Kortekaas and Dawson (2007), Morton et al. (2007).

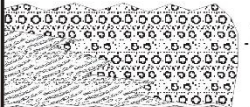
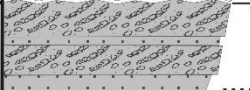
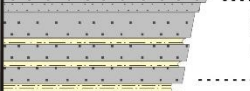
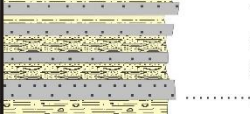

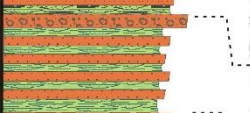
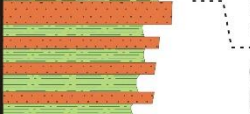
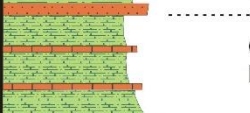
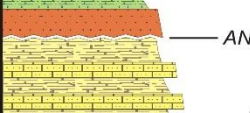
CHRONOZONE	FORMATION	BIOZONE	THICKNESS	LITHOLOGY	LITHOFACIES - LITHOLOGICAL EXPLANATIONS
U. Pleist.- Holocene	Fluvial (Qal)- coastal (Qac) sediments	-	-		Alluvial/colluvial/coastal settings: Terrestrial various-sized siliciclastic sediments / rocks Fluvial-dominated delta: red-brown cross-bedded conglomerate-sandstone alternation ANGULAR UNCONFORMITY
Lower Pleistocene	Manavgat Formation (Qm) "Tsunamite" member (Qmt)	NN19a - Gephyrocapsa caribbeanica zone	~ 400 m		Wave-dominated delta: type cross-bedded grey-brown conglomerate-sandstone alternation
					Shoreface/nearshore: yellowish-grey coarse-sandstone and laminated sandstone / sandy mudstone alternation
Lower Pliocene	Karpuzçay Formation (Nk)	NN15 - Reticulofenestra umbilica zone	~ 750 m		Transition between offshore shelf & lower shore-face: alternations of hummocky cross-stratified, brown-grey sandstone highly burrowed, brown-grey silty mudstones
					Tsunamite zone: Multigenetic pebbly mud matrix including recrystallized limestone (Mesozoic), conglomerate-sandstone (Neogene) cobble/blocks, marl blocks, reefal boulders/cobbles
					ANGULAR UNCONFORMITY - Late Pleistocene Erosional Surface (LPES) - Upper delta / tidal & coastal flat: alternations of red-brown sandstone and highly burrowed, wavy/flaser cross-bedded green-grey mudstones including coal lenses
Langhian - Serravalian	Geceleme Formation (Ng)	NN5 - Sphenolithus heteromorphus zone	~ 50 m		Clastic & carbonate shelf: Alternations of red-green limestone/calcarenite, blue-green marls
	Oymapınar Formation (No)	-	-		Carbonate shelf: Alternations of cream-light brown limestone/calcarenite, light brown marls
					CONFORMABLE BOUNDARY Neritic setting & carbonate shelf: Yellow-cream colored fossiliferous limestone

Figure 2. Generalized stratigraphic section taken from the Neogene–Quaternary sedimentary successions in the study area located at the Antalya-Manavgat Subbasin (SW Turkey).

Şekil 2. Antalya-Manavgat Alt Havzasında (GB Türkiye) yer alan çalışma alanındaki Neojen-Kuvaterner sedimanter istiflerden alınan genelleştirilmiş stratigrafik kesit.

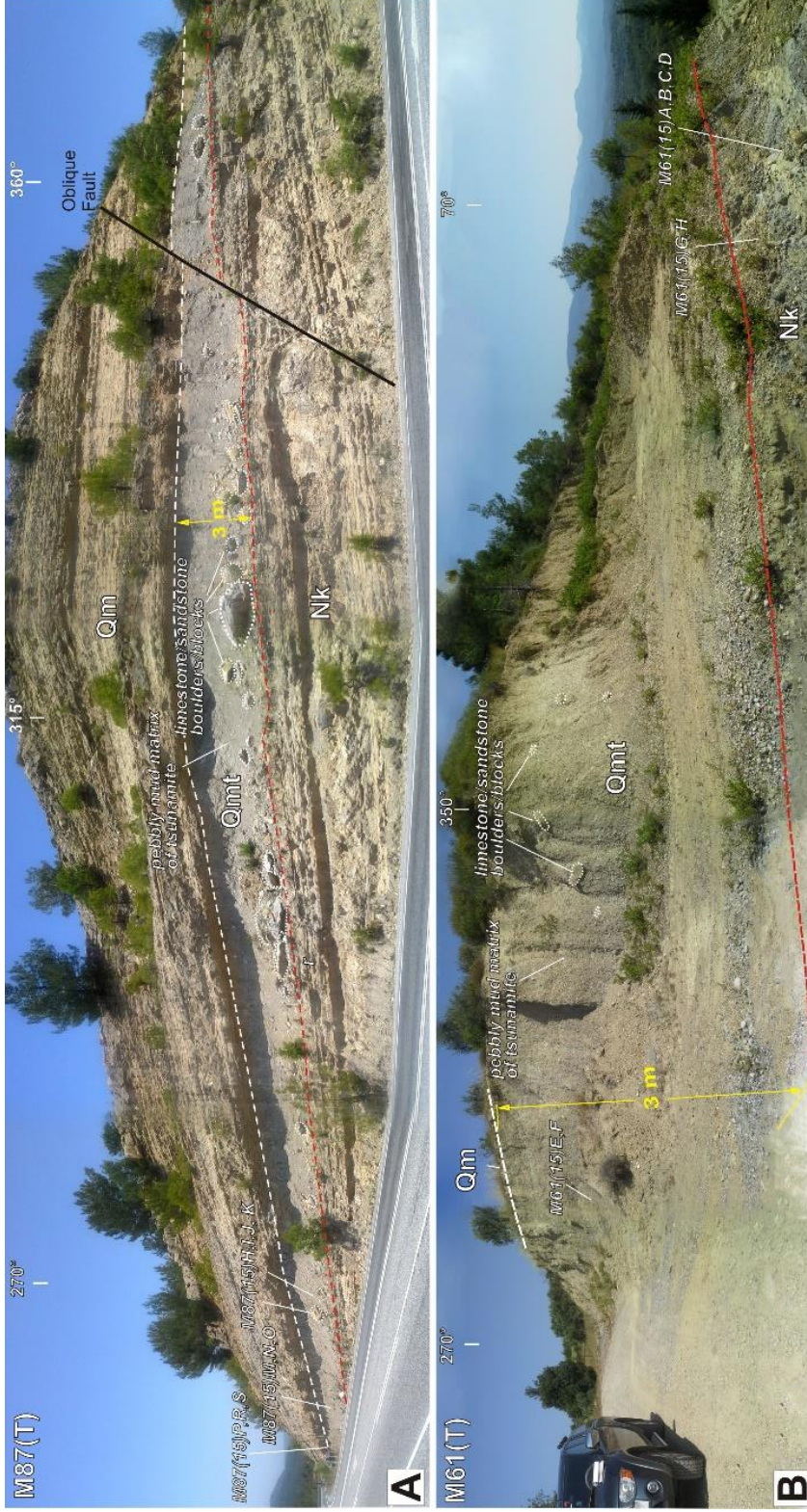


Figure 3. Field images of M61(T) and M87(T) investigation spots include the tsunamite levels between Plio-Quaternary sedimentary successions in the study area in the Manavgat Subbasin (SW Turkey): A–B) Boulderly muddy tsunamite deposits (Qmt) in the lower levels of the Manavgat Formation (Qm) covering to the Karpuzçay Formation (Nk) and rock sample places (A: M87(T), B: M61(T), white dashed lines: upper boundary of tsunamite, red dashed lines: tsunamite lower boundary/LPES).

Şekil 3. M61(T) ve M87(T) inceleme noktalarının saha görüntüleri, Manavgat Alt Havzası (GB Türkiye) çalışma alanındaki Pliyo-Kuvaterner sedimanter istifler arasındaki tsunamite seviyelerini içermektedir: A–B) Karpuzçay Formasyonu'nun (Nk) örtlen Manavgat Formasyonu'nun (Qm) alt seviyelerinde bulunan, bloktu çamurlu tsunamite çökelleri (Qmt) ve kayaç örnek yerleri (A: M87(T), B: M61(T), beyaz kesikli çizgiler: tsunamite üst sınırı, kırmızı kesikli çizgiler : tsunamite alt sınırı/LPES).stratigrafik referans kesitler.

Among its blocks of resistant rocks such as limestone, conglomerate and sandstone, in the form of boulders, cobbles and pebbles, buried here in the tsunami mud matrix, there are well-preserved cobble-sized blocks of blue-green marl belonging to the lower regular sequence (Karpuzçay Formation) and reefal blocks of boulder or cobble size, coming from synsedimentary (Quaternary) environments. The thickness of the paleotsunamite level was measured about 2-3 meters in the field. As illustrated in detail in the other article of the authors (to be published by Sagular and Yavuzlar-Özer in 2024), example field images of M61(15) and M87(15) observation spots for the tsunamite levels are shown in Figure 3. Upwards, regular sedimentary rock levels of the Manavgat Formation firstly include hummocky cross-stratified, brown-grey sandstones alternated with brown-grey silty mudstones including common burrows of transition between offshore shelf and lower shore-face. In the upper levels, there are yellowish-grey coarse-sandstone and laminated sandstone or sandy mudstone alternations representing shoreface/nearshore conditions. In the uppermost levels, the formation includes wave-dominated delta type cross-bedded grey-brown conglomerate-sandstone alternation. All defined sedimentologic and stratigraphic data indicate that the sedimentation of the formation which started with a short transgressive period, continued in regressive character. (4) Last stratigraphic sequence (Late Pleistocene-Holocene aged) is to be still at the beginning of a new stratigraphic-sedimentologic cycle. In the first levels, the sequence contains fluvial-dominated delta deposits comprising red-brown cross-bedded conglomerate-sandstone alternation. Upwards, terrestrial, various-sized siliciclastic sediments/rocks consisting of mainly sandy/pebbly clastics transitionally deposited in alluvial/colluvial/coastal settings are seen in the study area.

4. FIELD OBSERVATIONS AND STRATIGRAPHIC REFERENCE SECTIONS RELATED THE TSUNAMITE LEVELS

In the study, detailed observations and rock-sampling in the 10 investigation spots at the unconformable boundary between Lower Pliocene aged Karpuzçay Formation (the lower levels of the former Karpuzçay Formation) and Lower Pliocene aged Manavgat Formation (the upper levels of the previously-called Karpuzçay Formation) were made (see Fig. 1). Based on the data obtained from 10 investigation spots, 8 measured stratigraphic reference sections i.e. M2(T), M53(T)-1/M53(T)-2, M60(T)/M61(T) (Fig. 3), M87(T) (Fig. 3),

M91(T), M94(T), M96(T), M105(T) were designed (Fig. 4). Based on the comparison of the 8 stratigraphic reference sections, to observe similar sedimentologic and stratigraphic features in all sections are remarkable.

In the 4 sections (see Fig. 4) designed from the investigation spots M53(T)-1/M53(T)-2, M60(T)/M61(T), M91(T), M96(T) are seen alternations of red-yellow sandstones and blue-grey marls. These units indicate these levels of the Karpuzçay Formation might have been deposited on clastic/carbonate shelf and in regressive character. Above the Late Pliocene Erosional Surface (LPES), this unit is unconformably overlain by approximately 2 m thick tsunami sediments representing the Manavgat Formation and then continues as regular regressive shallow marine clastics of the same formation. North-eastern spots M87(T) and M105(T), south-western spots M53(T)-1 / M53(T)-2, M60 / M61 (15), M91 (15), M96 (16) have the same sequence. But their tsunamite levels reaching up to 3 meters are higher than the others. M2(T) and M94(T) numbered spots expose about 1,5 meters tsunamite zone in the most south-western of the study area. In addition, in a pebbly mud matrix of the tsunami deposits of the north-eastern spots M87(T) and M105(T) include commonly boulder/block-sized reworked rock material whereas in the south-western spots M53(T)-1 / M53(T)-2, M60(T), M61(T), M91(T), M96 (16), there are mainly block/cobble-sized reworked material and in the most south-western spots, usually cobble-sized reworked rock fragments are exhibited. In addition to an increase in the thickness of the tsunami deposits, from the spots of M2(T), M53(T)-1 / M53(T)-2 to M60(T) / M61(T), M94(T), M105(T) and then in the direction of M91 (T), M87 (T), M96 (T), from southwest to northeast, their grain sizes increase is observed.

5. BIOSTRATIGRAPHIC DATA

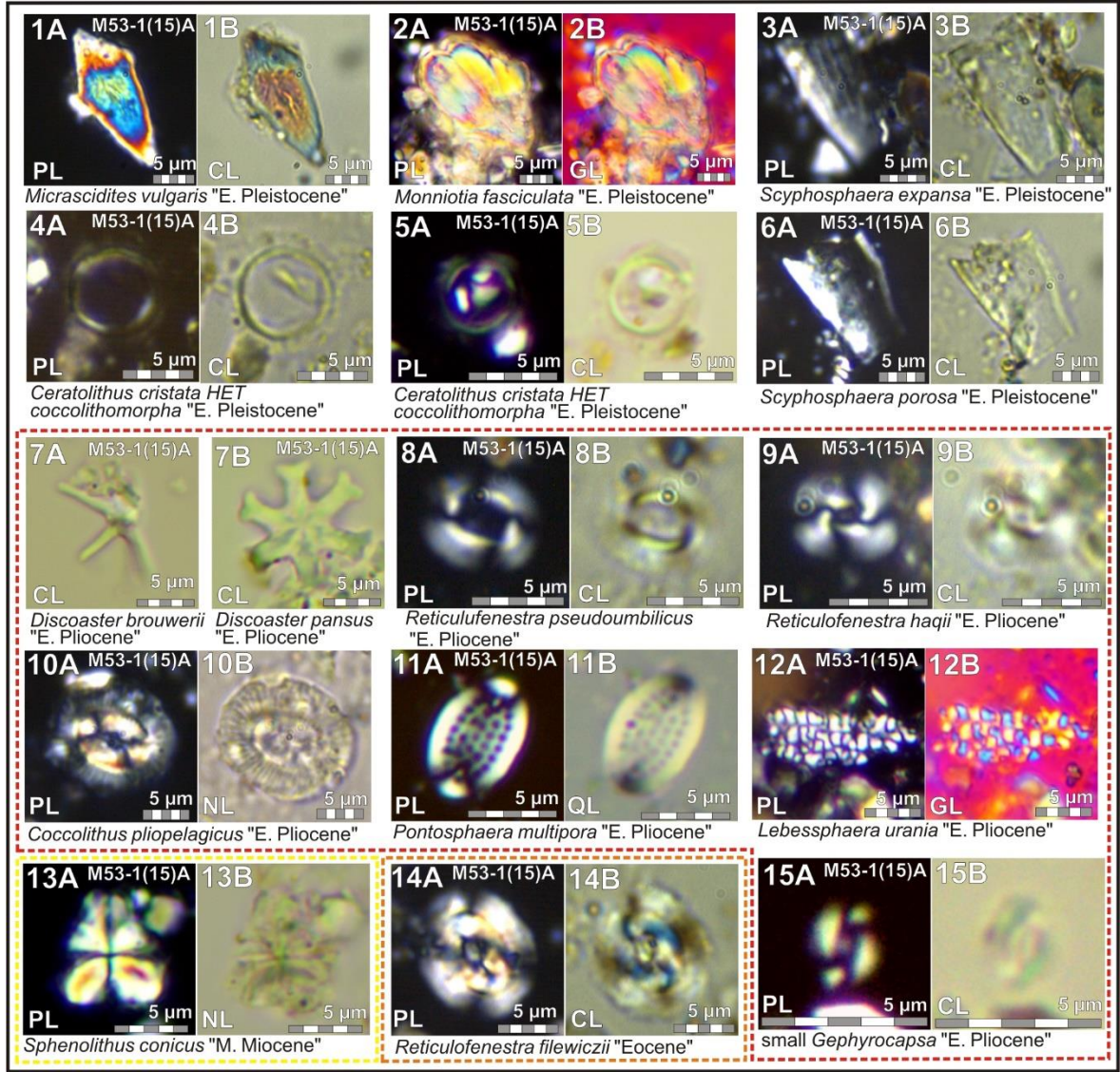
5.1. Originally Discriminated Nannofossil Data and Fossil Ascidian Spicule Records of the Karpuzçay-Manavgat Formations and Tsunami Deposits

47 various-sized clastic or calcareous rock samples derived from 10 investigation spots are shown in the 8 measured stratigraphic reference sections (see Fig. 4) and they have been examined for nannofossils by smear slides and thin sections under polarizing microscopes.

In these samples (M53-2(15)A, M61(15)B, M61(15)D, M87(15)H, M87(15)I, M91(15)B, and M96(16)B), there are Early Pliocene nannofossil assemblage of NN15 - Reticulofenestra pseudoumbilicus interval zone including synsedimentary data including *Calcidiscus* spp. [i.e. *C. leptoporus* (Murray & Blackman), *C. macintyreii* (Bukry & Bramlette), *C. tropicus* Kamptner], *Calciosolenia murrayi* Gran, *Coccolithus pliipelagicus* Wise, *Coronosphaera mediterranea* (Lohmann), *Discoaster* spp. [i.e. *D. asymmetricus* Gartner, *D. brouweri* Tan, *D. pansus* (Bukry & Percival), *D. pentaradiatus* Tan, *D. pliostellulus* Browning & Bergen, *D. surculus* Martini & Bramlette, *D. tamalis* Kamptner, *D. variabilis* Martini & Bramlette], *Helicosphaera* spp. [i.e. *H. intermedia* Martini, *H. kamptneri* Hay & Mohler, *H. selli* (Bukry and Bramlette)], *Pseudoemiliana ovata* (Kamptner), *Reticulofenestra* spp. [i.e. *R. haqii* Backman, *R. minuta* Roth, *R. minutula* (Gartner), *R. pseudoumbilicus* (Gartner)], *Scyphosphaera* spp. (i.e. *S. apsteinii* Lohmann, *S. canescens* Kamptner, *S. intermedia* Deflandre, *S. pulcherrima* Deflandre), *Sphenolithus* spp. [i.e. *S. abies* Deflandre, *S. neoabies* Bukry & Bramlette, *Sphenaster metula* (Wilcoxon)], *Syracosphaera pulchra* Lohmann, *Umbilicosphaera* spp. [i.e. *U. jafarii* Müller, *U. rotula* Kamptner, *U. sibogae* (Weber-van Bosse)]. In addition, some dinoflagellate cysts such as *Calciodinellum albatrosianum* (Kamptner) (see Fig 6: 15a-b), *Lebessphaera urania* Meier, Janofske & Willems were determined.

The resembling data were obtained from the blue-grey marl block/cobbles placed within the tsunamite matrix. Synsedimentary nannofossil quantities show a distribution mainly varying between 46–100% in the samples (see Table 2). The resembling data in 43–98% proportions were obtained from 9 samples of the blue-grey marl blocks/cobbles (M2(15)B, M60(15)B, M61(15)E, M91(15)G, M94(16)D, M105(16)B, M105(16)D, M105(16)E, M105(16)F) participated into the tsunamite matrix (see Table 2).

In addition, in these samples of the Karpuzçay Formation, reworked nannofossil data of Early-Middle Miocene, Eocene, Cretaceous-Paleocene nannofossil assemblages are described. *Calcidiscus premacintyreii* Theodoridis, *Coccolithus miopelagicus* Bukry, *Cyclicargolithus* spp. (i.e. *C. abisectus* (Müller), *C. floridanus* (Roth & Hay)), *Dictyococcites bisectus* (Hay, Mohler and Wide), *Discoaster druggii* Bramlette & Wilcoxon, *Helicosphaera* spp. (i.e. *H. recta* (Haq), *H. stalis* Theodoridis, *H. walbersdorfensis* Muller), *Scyphosphaera recurvata* Deflandre, *Sphenolithus* spp. (i.e. *S. belemnos* Bramlette and Wilcoxon, *S. compactus* Backman, *S. conicus* Bukry, *S. dissimilis* Bukry and Percival, *S. heteromorphus* Deflandre, *S. moriformis* (Brönnimann &



EXPLANATIONS (for Figures 5-9):

- E. Pleistocene syndesimentary nannofossils in tsunamite matrix or mudstone of Manavgat Fm.)
 - E. Pleistocene nannofossils in intraclast
 - E. Pliocene syndesimentary nannofossils in sedimentary rocks of Karpuzçay Fm.
 - E. Pliocene syndesimentary nannofossils in blue marl block
 - E. Pliocene (reworked) nannofossils
 - Miocene syndesimentary nannofossils in brown-yellow marl block
 - M. Miocene (reworked) nannofossils
 - Eocene (reworked) nannofossils
 - Cretaceous-Paleocene (reworked) nannofossils
- (Int): Intraclast sample, (Tm): Tsunamite matrix
 - PL Polarized Light
 - NL Normal Light
 - CL Phase Contrast
 - QL With Quartz Wedge
 - GL With Gypsum Wedge

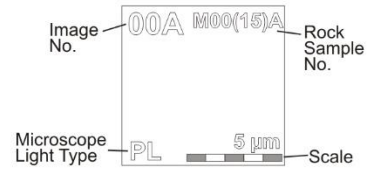


Figure 5. Syndesimentary and reworked nannofossil species in the mudstone sample M53-1(15)A of the Manavgat Formation (regular sedimentary sequence).

Şekil 5. Manavgat Formasyonu'na ait M53-1(15)A çamurtaşı örneğindeki çökelimle eşyaşlı ve yeniden çökeltilmiş nannofosil türleri (düzenli sedimanter istif).

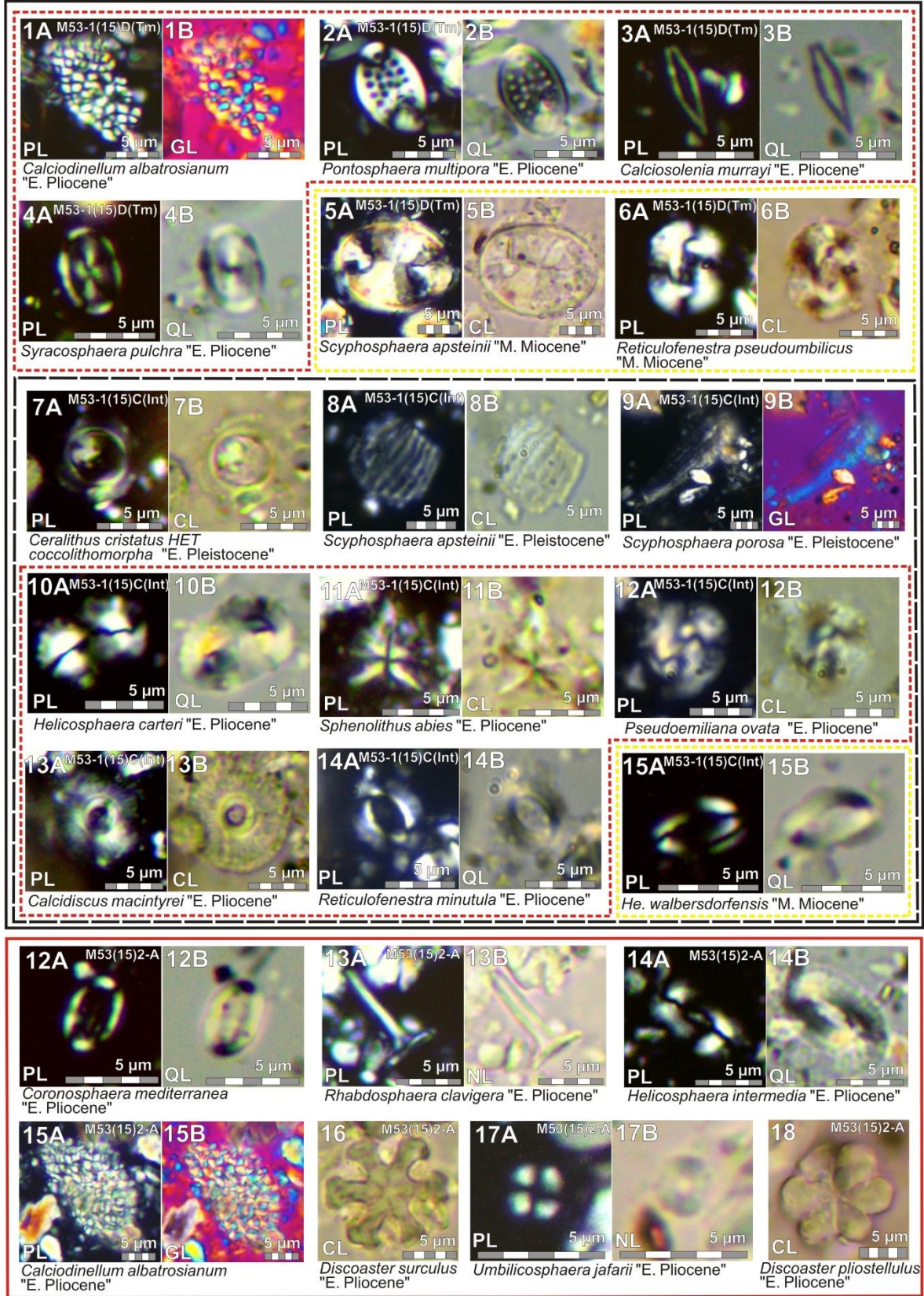


Figure 6. Synsedimentary and reworked nannofossil species in the (pebbly) mud matrix (M53-1(15)D), mud intraclast (M53-1(15)C) samples in the tsunami deposits of the Manavgat Formation and blue marl (M53-2(15)A) of the Karpuzçay Formation (Explanations in Fig. 5).

Şekil 6: Manavgat Formasyonu'nun tsunami çökellerindeki (çakıllı) çamur matriks (M53-1(15)D), çamur intraklast (M53-1(15)C) ve Karpuzçay Formasyonu'nun mavi marn (M53-2(15)A) örneklerindeki çökelimle eşyaşlı ve yeniden çökeltilmiş nannofosil türleri (Açıklamalar Şekil 5'te).

Stradner)), *Triquetrorhabdulus carinatus* Martini, *Umbilicosphaera jafarii* Müller are nannofossil data reworked from Early-Middle Miocene aged sedimentary rocks such as Middle Miocene aged Geceleme Formation north of the region or the older Lower Miocene rocks around Isparta.

Nannofossil assemblage of the Geceleme Formation indicates NN5 - *Sphenolithus heteromorphus* zone whereas an Aquitanian-Burdigalian assemblage represents nannofossil data of NN1 to NN3 biozones determined around Isparta (Görmüş et al, 2001; Hepdeniz and Sagular, 2009; Yavuzlar and Sagular, 2018). With a few species and numbers, *Chiasmolithus solitus* (Bramlette and Sullivan), *Coccolithus pelagicus* (Wallich), *Discoaster deflandrei* Bramlette & Riedel, *Ericsonia* spp. (i.e. *E. formosa* (Kamptner), *E. ovalis* Black), *Reticulofenestra* spp. (i.e. *R. dictyoda* (Deflandre), *R. hillae* Bukry & Percival, *R. reticulata* (Gartner & Smith), *R. stavensis* (Levin & Joerger)), *Sphenolithus moriformis* (Brönnimann & Stradner), *Toweius* spp. (i.e. *T. callosus* Perch-Nielsen, *T. magnicrassus* (Bukry)), *Zygrhablithus bijugatus* (Deflandre) are found as 1-7 % Eocene nannofossil assemblage in blue-grey marl samples of the Karpuzçay Formation.

Besides Upper Cretaceous (reworked) nannofossil data including *Watznaueria barnesiae* (Black in Black & Barnes), *Zeughrabdotos sigmoides* (Bramlette & Sullivan) and dinoflagellate cysts such as *Cervisiella saxea* (Stradner) are seen in the samples of Karpuzçay Formation or various marl blocks/cobbles in the tsunamite matrix. In addition, there are light brown-cream colored marly limestone, marl cobbles/blocks [samples: M87(15)C(BI), M87(15)F(BI), M94(16)C(BI)] including Lower-Middle Miocene nannofossils (see Table 3) such as *Calcidiscus premacintyreii* Theodoridis, *Cyclicargolithus floridanus* (Roth & Hay), *Dictyococcites* spp. (*D. productus* (Kamptner), *D. scripsae* Bukry & Percival), *Helicosphaera* spp. (i.e. *H. carteri* (Wallich), *H. walbersdorfensis* Muller), *Reticulofenestra* spp. (i.e. *R. haqii* Backman, *R. minuta* Roth, *R. pseudoumbilica* (Gartner)), *Scyphosphaera* spp. (i.e. *S. apsteinii* Lohmann, *S. recurvata* Deflandre), *Sphenolithus* spp. (i.e. *S. compactus* Backman, *S. conicus* Bukry, *S. heteromorphus* Deflandre, *S. moriformis* (Brönnimann & Stradner)), *Umbilicosphaera jafarii* Müller representing NN5 zone. It is concluded that these pebbles/cobbles (rock samples) were

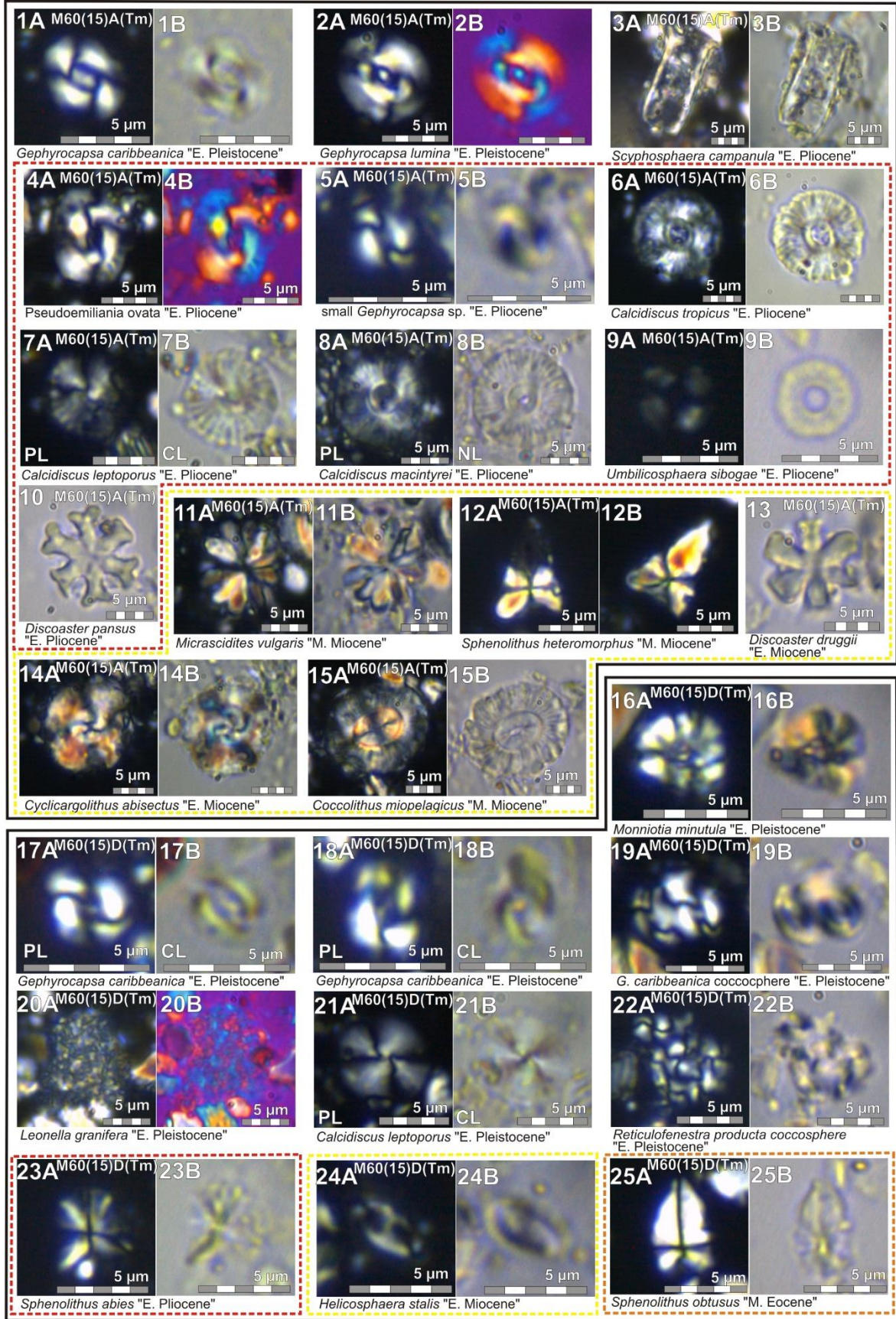


Figure 7. Synsedimentary and reworked nannofossil species in the (pebbly) mud matrix (M60(15)A and M60(15)D) samples in the tsunami deposits of the Manavgat Formation. (Explanations in Fig. 5).

Şekil 7. Manavgat Formasyonu'nun tsunami tortullarındaki (M60(15)A ve M60(15)D) örneklerinin (çakıllı) çamur matrisindeki çökelimle eşyaşlı ve yeniden çökeltilmiş nannofosil türleri. (Açıklamalar Şekil 5'te).

probably transported from carbonate rocks of the Geceleme Formation placed at the bottom. As illustrated in Figures 6, 8, 12, 13, Lower Pleistocene nannofossils of the Karpuzçay Formation were determined in a number of the blue-grey marl samples [i.e. M53-2(15)A in Fig. 6, M61(15)B in Fig. 8, M96(16)B in Fig. 12] and in the blue-grey marl cobble/blocks of the tsunamite [i.e. B61(15)E(BI) and M61(15)G(BI) in Fig. 8, M105(16)B(BI) in Fig. 13].

As illustrated in Figure 5–13, reworked Lower Pliocene nannofossils are between 36–87% in the Lower Pleistocene regular (rhythmic shallow marine clastic) sedimentary rock samples of the Manavgat Formation [i.e. M53-1(15)A in Fig. 5; M87(15)O, M87(15)S in Fig. 9; M94(16)B in Fig. 11; M96(16)F in Fig. 12] and are between 71–75% in the gray mudstone intraclast samples [i.e. M53-1(15)C(Int) in Fig. 6] and are between 51–98% in the irregular (tsunamite) sedimentary rock samples [i.e. M53-1(15)D(Tm) in Fig. 6; M60(15)A(Tm), M60(15)D(Tm) in Fig. 7; M94(16)A(Tm) in Fig. 10; M96(16)C(Tm) in Fig. 12; M105(16)A(Tm) in Fig. 13]. In addition, an abundant reworked Lower-Middle Miocene nannofossil (and dinoflagellate cysts) data between 1–55% (see Table 3, Figs. 5-13) including *Coccolithus miopelagicus* Bukry, *Cyclicargolithus* spp. [i.e. *C. abisectus* (Müller), *C. floridanus* (Roth & Hay)], *Dictyococcites bisectus* (Hay, Mohler and Wide), *Discoaster druggii* Bramlette & Wilcoxon, *Helicosphaera* spp. (i.e. *H. stalis* Theodoridis, *H. walbersdorfensis* Müller), *Micrantholithus flos Deflandre*, *Pontosphaera desuetoidea* Bartol, *Reticulofenestra* spp. [i.e. *R. haqii* Backman, *R. gelida* (Geitzenauer), *R. lockeri* Müller, *R. minuta* Roth, *R. pseudoumbilica* (Gartner)], *Scyphosphaera* spp. [i.e. *S. apsteinii* Lohmann, *S. brevis* Varol, *S. lagena* Kamptner, *S. recurvata* Deflandre, *S. tubifera* Kamptner], *Sphenolithus* spp. [i.e. *S. belemnus* Bramlette and Wilcoxon, *S. compactus* Backman, *S. conicus* Bukry, *S. disbelemnus* Fornaciari and Rio, *S. dissimilis* Bukry and Percival, *S. heteromorphus* Deflandre, *S. moriformis* (Brönnimann & Stradner)], *Thoracosphaera heimii* (Lohmann), *Triquetrorhabdulus carinatus* Martini, *Umbilicosphaera*

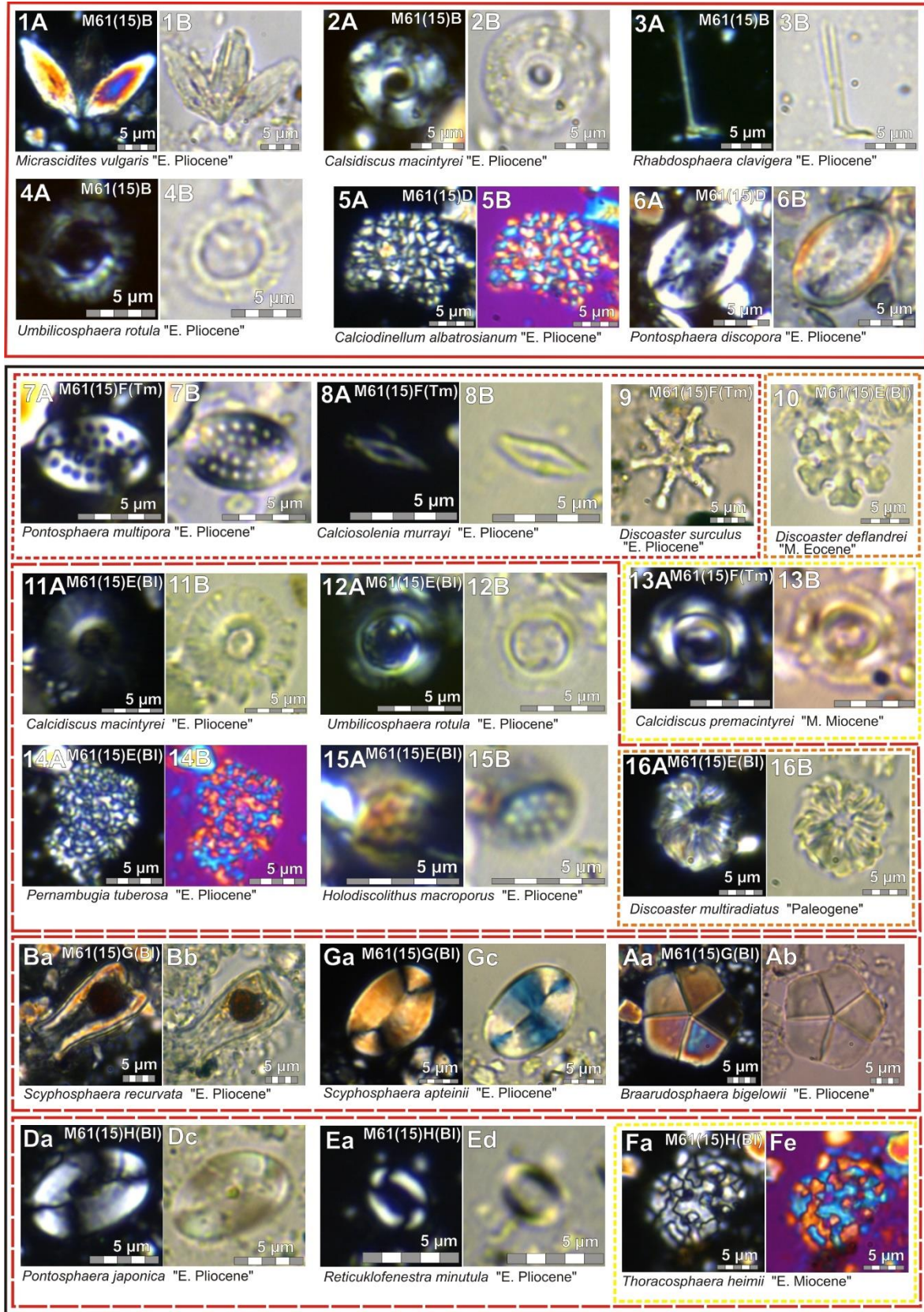


Figure 8. Syngedimentary and reworked nannofossil species in the blue-grey marl (M61(15)B and M61(1)D) samples of the Karpuzçay Formation; (pebbly) mud matrix sample (M61(15)A) and

blue-gray marl blocks (M61(15)E, M61(15)G and M61(15)H) in the tsunami deposits of the Manavgat Formation. (Explanations in Fig. 5).

Şekil 8. Karpuzçay Formasyonu'nun mavi-gri marn (M61(15)B ve M61(1)D) örnekleri; Manavgat Formasyonu'nun tsunami çökellerindeki (çakıllı) çamur matris örneği (M61(15)A) ve mavi-gri marn blokları (M61(15)E, M61(15)G ve M61(15)H) örneklerindeki çökelimle eşyaşlı ve yeniden çökeltilmiş nannofosil türleri. (Açıklamalar Şekil 5'te).

jafarii Müller. It was determined that a number of reworked Middle Eocene nannofossil data between 1–32% (see Table 4, Figs. 5–13) in the most of the rock samples of the Karpuzçay and Manavgat Formations. *Chiasmolithus solitus* (Bramlette and Sullivan), *Cruciplacolithus latipons* Romein, *Dictyococcities hesslandii* (Haq), *Discoaster* spp. (i.e. *D. backmanii* Agnini et al., *D. deflandrei* Bramlette & Riedel, *D. multiradiatus* Bramlette & Riedel), *Ericsonia* spp. [i.e. *E. formosa* (Kamptner), *E. ovalis* Black, *E. subpertusa* Hay & Mohler], *Pontosphaera* cf. *pectinata* (Bramlette & Sullivan), *Pseudotriquetrorhabdulus inversus* (Bukry & Bramlette), *Reticulofenestra* spp. [i.e. *R. dictyoda* (Deflandre in Deflandre & Fert), *R. hillae* Bukry & Percival, *R. reticulata* (Gartner & Smith), *R. stavensis* (Levin & Joerger)], *Sphenolithus moriformis* (Brönnimann & Stradner), *Toweius* spp. [i.e. *T. callosus* Perch-Nielsen, *T. magnicrassus* (Bukry)], *Zygrhablithus bijugatus* (Deflandre in Deflandre & Fert) are among of these data. In determinations were come across reworked Upper Cretaceous-Danian nannofossil and dinoflagellate species between 1-32% (see Table 4, Figs. 5–14) in some of the rock samples of the Karpuzçay and Manavgat Formations. *Calculites obscurus* (Deflandre), *Cribrosphaerella ehrenbergii* (Arkhangelsky), *Eiffellithus* spp. [i.e. *E. eximius* (Stover), *E. gorkae* Reinhardt, *E. turriseiffelii* (Deflandre)], *Lithraphidites quadratus* Bramlette & Martini, *Placozygus fibuliformis* (Reinhardt), *Retecapsa crenulata* (Bramlette & Martini), *Watznaueria barnesiae* (Black in Black & Barnes), *Zeugrhabdotus sigmoides* (Bramlette & Sullivan), *Cervisiella* spp. [I.e. *C. operculata* (Bramlette & Martini), *C. saxea* (Stradner)] are among these data.

In the whole examined rock samples gathered from the regular sedimentary rock alternations as rhythmic shallow marine clastics [i.e. M53-1(15)A, M60(15)E, M87(15)R, M87(15)S, M94(16)B, M96(16)F] and irregular sedimentary levels “tsunami deposits” [i.e. M2(15)A(Tm), M53-1(15)D(Tm), M60(15)A(Tm), M61(15)F(Tm), M87(15)L(Tm), M87(15)N(Tm), M87(15)O(Tm), M91(15)F(Tm), M94(16)A(Tm), M96(16)C(Tm),

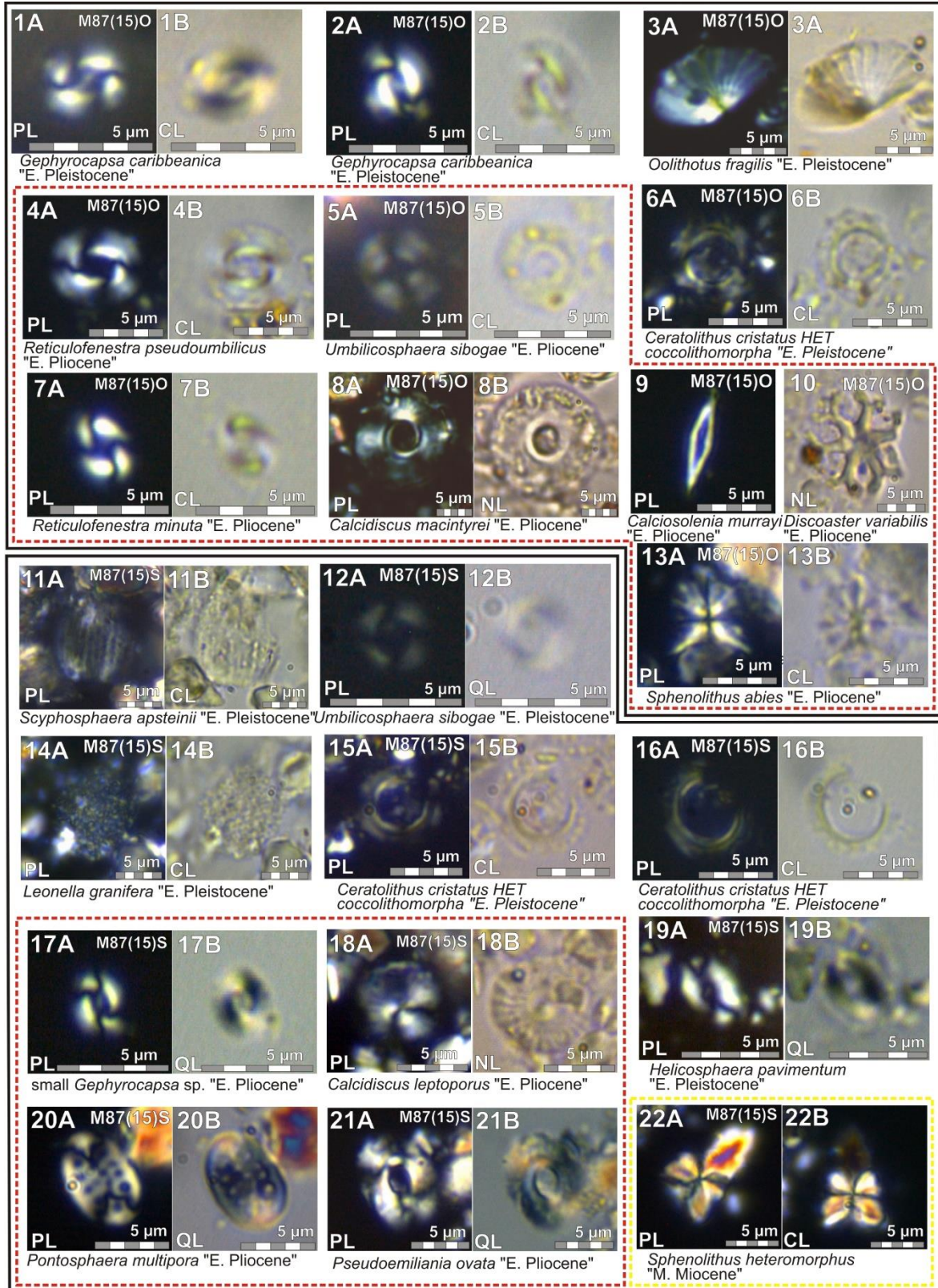


Figure 9. Synsedimentary and reworked nannofossil species in the (pebbly) mud matrix (M87(15)O and M87(15)S) samples in the tsunami deposits of the Manavgat Formation. (Explanations in Fig. 5).

Şekil 9. Manavgat Formasyonu'nun tsunami çökellerinin (çakıllı) çamur matrisindeki (M87(15)O ve M87(15)S) çökelimle eşyaşlı ve yeniden çökeltilmiş nannofosil türleri. (Açıklamalar Şekil 5'te).

M105(16)A(Tm)] of the Manavgat Formation, an abundance of reworked nannofossil data are seen to hide absolute synsedimentary data (see Tables 1-5, Figs. 5–14).

As some of them illustrated in Figs. 5–13, *Calcidiscus leptoporus* (Murray & Blackman), *Ceratolithus cristatus* HET coccolithomorpha (Lecal-Schlauder), medium *Gephyrocapsa* spp. (*G. caribbeanica* Boudreaux & Hay, *G. lumina* Bukry) and small *Gephyrocapsa* spp., *Helicosphaera* spp. (i.e. *H. carteri* (Wallich), *H. hyalina* Gaarder, *H. pavementum*), *Oolithotus fragilis* (Lohmann), *Pontosphaera* spp. [i.e. *P. japonica* (Takayama), *P. scutellum* Kamptner], *Pseudoemiliana lacunosa* (Kamptner), *Scyphosphaera* spp. (i.e. *S. apsteinii* Lohmann, *S. expansa* Bukry & Percival, *S. porosa* Kamptner), *Syracosphaera pulchra* Lohmann, *Umbellosphaera tenuis* (Kamptner), *Umbilicosphaera* spp. [i.e. *U. foliosa* Kamptner, *U. sibogae* (Weber-van Bosse)] are synsedimentary nannofossil data (see Table 1).

They have been seen generally between 3–14% (Fig. 15) in shallow marine mudstone-sandstone alternation of the Manavgat Formation, not exceeding 47% in some extreme samples, as well. In addition in the samples were determined a few dinoflagellate cyst species such as *Leonella granifera* (Fütterer), *Pernambugia tuberosa* (Kamptner) (see Table 1). Some samples such as M53-1(15)A, M60(15)E, M87(15)R, M87(15)S, M94(16)B, M96(16)F of regular sedimentary rock levels of the Manavgat Formation include total 53–97% reworked nannofossil data resembling those of the Karpuzçay and Geceleme Formations located at the region. These data comprise 36–87% Lower Pliocene nannofossil assemblage reworked from the Karpuzçay Formation rocks and 3–31% Lower-Middle Miocene assemblage reworked from the Geceleme Formation or the older rocks in and around the Isparta Angle (see Table 1–5, Figs. 5–14). Besides in the same samples, there are also 1–3% Middle Eocene and 1–6% Upper Cretaceous-Danian nannofossils reworked from the Eocene and Cretaceous/Paleogene transition rock sources located at the north of the Isparta Angle. The samples [M2(15)A(Tm), M53-1(15)D(Tm), M60(15)A(Tm), M60(15)D(Tm), M61(15)F(Tm), M87(15)L(Tm), M87(15)N(Tm), M87(15)O(Tm), M91(15)F(Tm), M94(16)A(Tm), M96(16)C(Tm), M105(16)A(Tm)] of pebbly mud matrix of the tsunamite deposits include mainly reworked nannofossil data

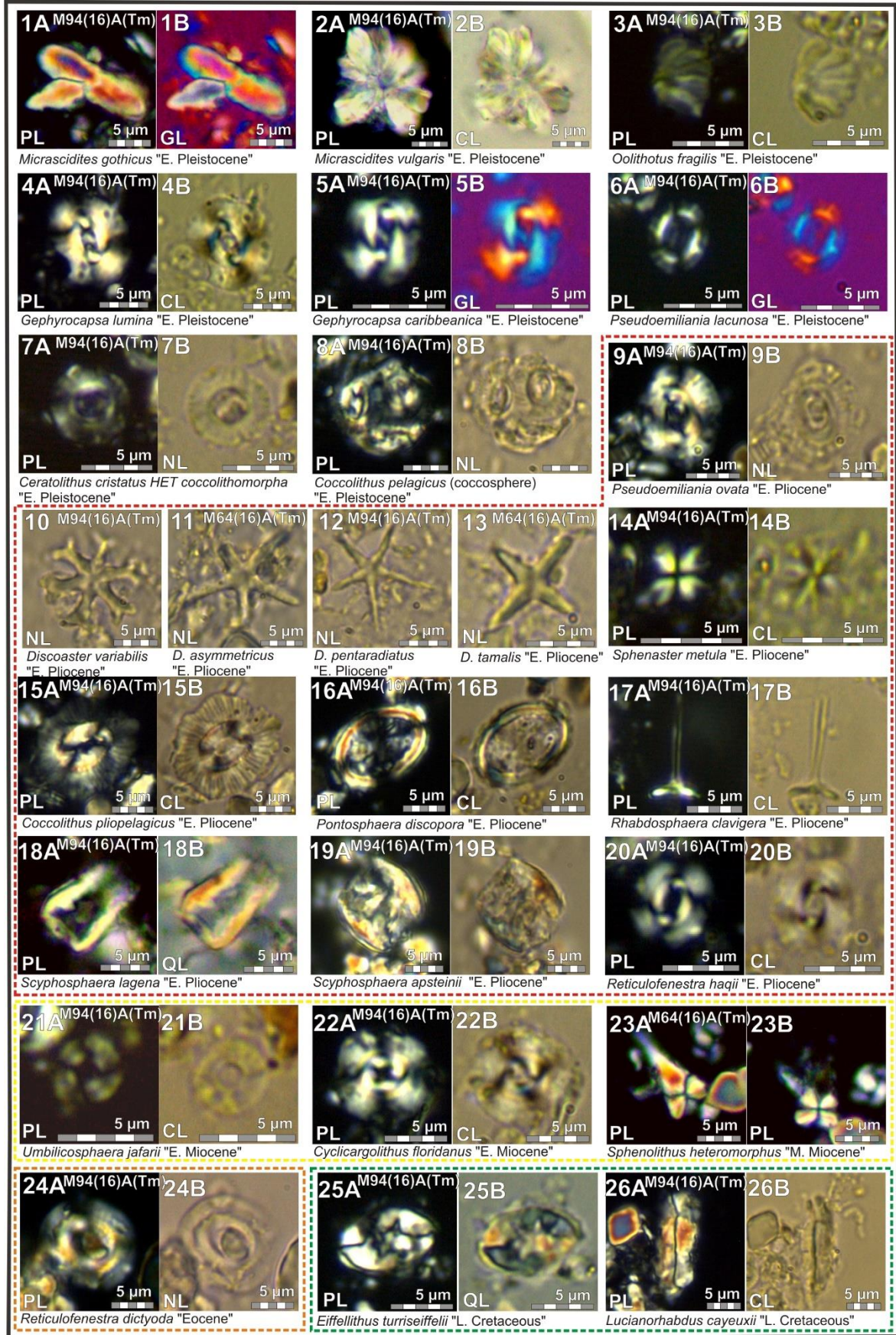


Figure 10. Syngsedimentary and reworked nannofossil species in the (pebbly) mud matrix (M94(16)A) sample in the tsunami deposits of the Manavgat Formation. (Explanations in Fig. 5)

Şekil 10: Manavgat Formasyonu'nun tsunami çökellerinde (çakıllı) çamur matris (M94(16)A) örneğindeki çökelimle eşyaşlı ve yeniden çökeltilmiş nannofosil türleri. (Açıklamalar Şekil 5'te).

derived from the Karpuzçay and Geceleme Formations between 60–96% while syngsedimentary fossils are in 4–40% (see Tab. 1–5, Fig. 5–14).

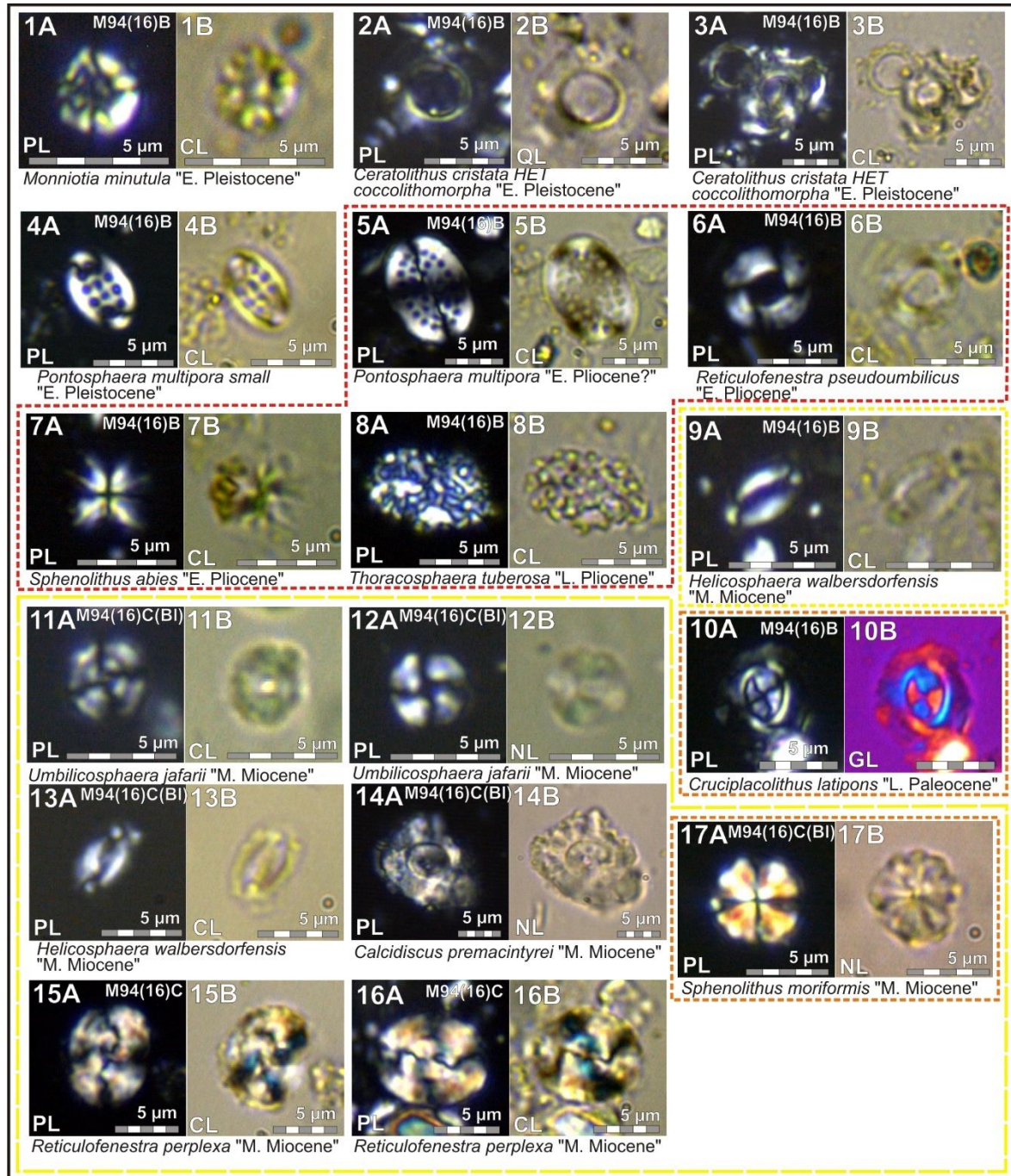


Figure 11. Synsedimentary and reworked nannofossil species in the (pebbly) mud matrix (M94(16)A) sample in the tsunami deposits of the Manavgat Formation) and light brown marl block (M94(16)C) from the carbonate shelf deposits of the Geceleme Formation. (Explanations in Fig. 5).

Şekil 11. Manavgat Formasyonu'nun tsunami tortullarından (çakıllı) çamur matriks (M94(16)A) örneği ve Geceleme Formasyonu'nun karbonat şelf çökellerinden açık kahverengi marn bloğu (M94(16)C) içindeki çökelimle eşyaşlı ve yeniden çökeltilmiş nannofossil türleri. (Açıklamalar Şekil 5'te).

Rarely encountered fossil ascidian spicules such as *Micrascidites vulgaris* Deflandre, Deflandre-Rigaud were determined in the sample blue-grey marl [M61(15)B] of Karpuzçay Formation (see Tab. 6, Fig. 8). Additionally, more fossil ascidian spicules such as *Bonetia acuta* Varol & Houghton, *Bonetia brevis* Varol & Houghton, *Micrascidites gothicus* Sagular in Sagular et al., *Micrascidites vulgaris* Deflandre, Deflandre-Rigaud, *Monniotia aciformis* Varol & Houghton, *Monniotia fasciculata* Varol & Houghton, *Monniotia minutula* Sagular in Sagular et al., *Riguadia multiradiata* Varol & Houghton were recorded in the grey-brown mudstone and yellow-brown sandstone samples [M53-1(15)B(Ts), M60(15)E, M96(16)F] above the tsunamite of the Manavgat Formation and/or pebbly grey-green mud matrix samples [M2(15)A(Tm), M53-1(15)A(Tm), M60(15)A(Tm), M94(16)A(Tm), M96(16)C(Tm)], in addition some muddy rip-up clasts [i.e. M105(16)D(Int)] of the tsunamite matrix (see Tab. 6, Figs. 5, 7, 10, 11).

5.2. Biochronozons of the Karpuzçay and Manavgat Formations Based on Nannofossil Data and Their Stratigraphic-Sedimentologic Origins

In this study, Karpuzçay Formation which was defined a Miocene marine deposition by Monod, O., 1977, Poisson, 1977; Akay et al., 1985), Şenel (1997), Flecker et al. (1998), Robertson (1998), Karabıyıköğlü et al. (2000), Robertson et al. (2003), Flecker et al. (2005), Karabıyıköğlü et al. (2005), Çiner et al. (2008) has been divided into two sequences: Lower Pliocene Karpuzçay Formation and Lower Pleistocene Manavgat Formation based on discriminated nannofossils according to their sedimentary-stratigraphic origins (see Tabs. 1–5). There are several origins of nannofossil particles participating in forming of a marine sedimentary rock: 1) "Synsedimentary nannofossils" collapsed as living coccospheres or individual nannoliths spilling out from these

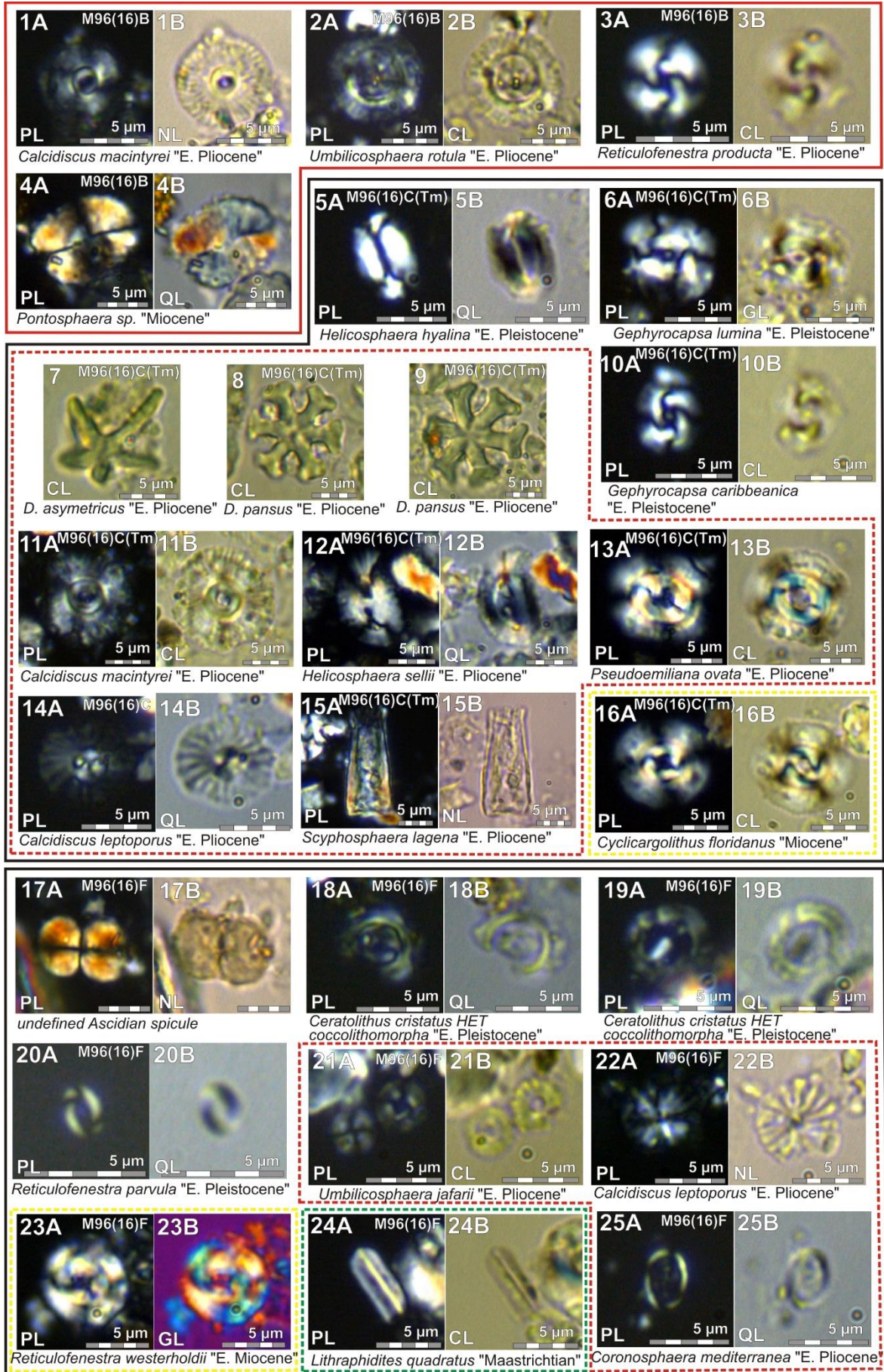


Figure 12. Synsedimentary and reworked nannofossil species in blue marl (M96(16)B) of the Karpuzçay Formation, in the (pebbly) mud matrix (M96(16)C) sample of the tsunamite and in the mudstone (M96(16)F) sample collected from regular clastic rocks of the Manavgat Formation (Explanations in Fig. 5).

Şekil 12. Karpuzçay Formasyonu'nun mavi marnında (M96(16)B), tsunamitin (çakıllı) çamur matriksinde (M96(16)C) ve Manavgat Formasyonu'nun düzenli kırıntılı kayalarından toplanan çamurtaşı (M96(16)F) örneğinde çökelimle eşyaşlı ve yeniden çökeltilmiş nannofosil türleri (Açıklamalar Şekil 5'te).

nannoplanktons, 2) "reworked nannofossils" redeposited nanno-bearing extrabasinal sedimentary rock sequences, 3) "removed nannofossils" translocated close-aged nanno-bearing intrabasinal sediments by means of marine wave¤ts, submarine slump&slides, and bottom erosions. In the study related to the discussive Neogene marine sequences located at the Manavgat Subbasin, "synsedimentary" and "reworked" nannofossil assemblages were discriminated according to sedimentary and stratigraphic origins of various-aged nannofossil assemblages and marine depositions.

Five essential sedimentary sequences and their nannoplankton assemblages were defined in the study: 1) Firstly Upper Cretaceous – Danian carbonates and clastics, having mainly K/Pg transition nannofossil data; 2) Middle Eocene carbonates and clastics, with the same-aged nannofossils; 3) Middle Miocene marine carbonates and clastics of the Geceleme Formation, including nannofossil assemblage of NN5 zone; 4) Lower Pliocene marine carbonates and clastics of (re-defined) Karpuzçay Formation, with nannofossil assemblage of NN15 zone; 5) Lower Pleistocene marine clastics of (first suggested) Manavgat Formation, containing nannofossils representing NN19 zone/CN13b subzone (see Figs. 15–16). Upper Cretaceous-Lutetian nannofossil data determined in the Karpuzçay and Manavgat Formations were probably reworked from the Cretaceous–Paleocene transition sequence on the Antalya Nappes north of Isparta Angle and Lutetian clastic sequence in the Akseki Region (Şenel, 1997), since pre-Miocene sedimentary rocks could not be observed on the ophiolitic basement of the Manavgat Subbasin in the former studies (i.e. Glover and Robertson, 1998) around Isparta Angle. 12 Upper Cretaceous-Danian nannofossils including *Eiffellithus* spp. [*E. gorkae* Reinhardt, *E. turriseiffelii* (Deflandre)], *Lithraphidites quadratus* Bramlette & Martini, *Watznaueria barnesiae* (Black in Black & Barnes), *Zeughrabdotus sigmoides* (Bramlette & Sullivan) and dinoflagellate cysts such as *Cervisiella operculata* (Bramlette & Martini), *Cervisiella saxea* (Stradner) are seen to have

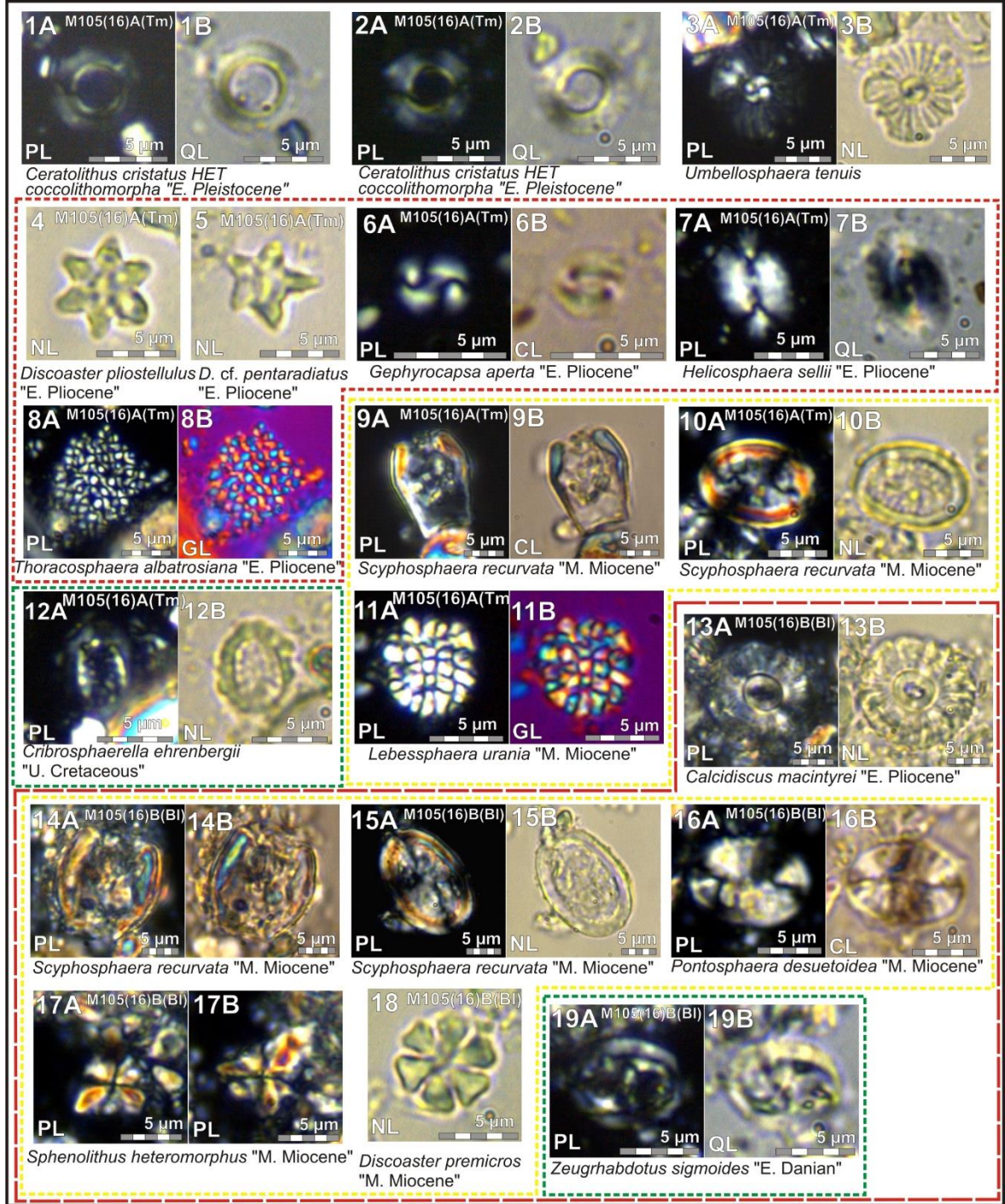


Figure 13. Synsedimentary and reworked nannofossil species in the (pebbly) mud matrix (M105(16)A) sample of tsunami deposits (Manavgat Formation) and blue marl block sample (M105(16)B) reworked from Karpuzçay Formation (Explanations in Fig. 5).

Şekil 13. Tsunami tortularına (Manavgat Formasyonu) ait (çakıllı) çamur matris (M105(16)A) örneği ve Karpuzçay Formasyonu'ndan yeniden çökeltilmiş mavi marn bloğu örneğindeki (M105(16)B) çökelle eşyaşlı ve yeniden çökeltilmiş nannofosil türleri (Açıklamalar Şekil 5'te).

participated into Geceleme, Karpuzçay and Manavgat Formation sedimentations as “reworked nannofossil” sedimentary particles (Tables 5, Figure15).

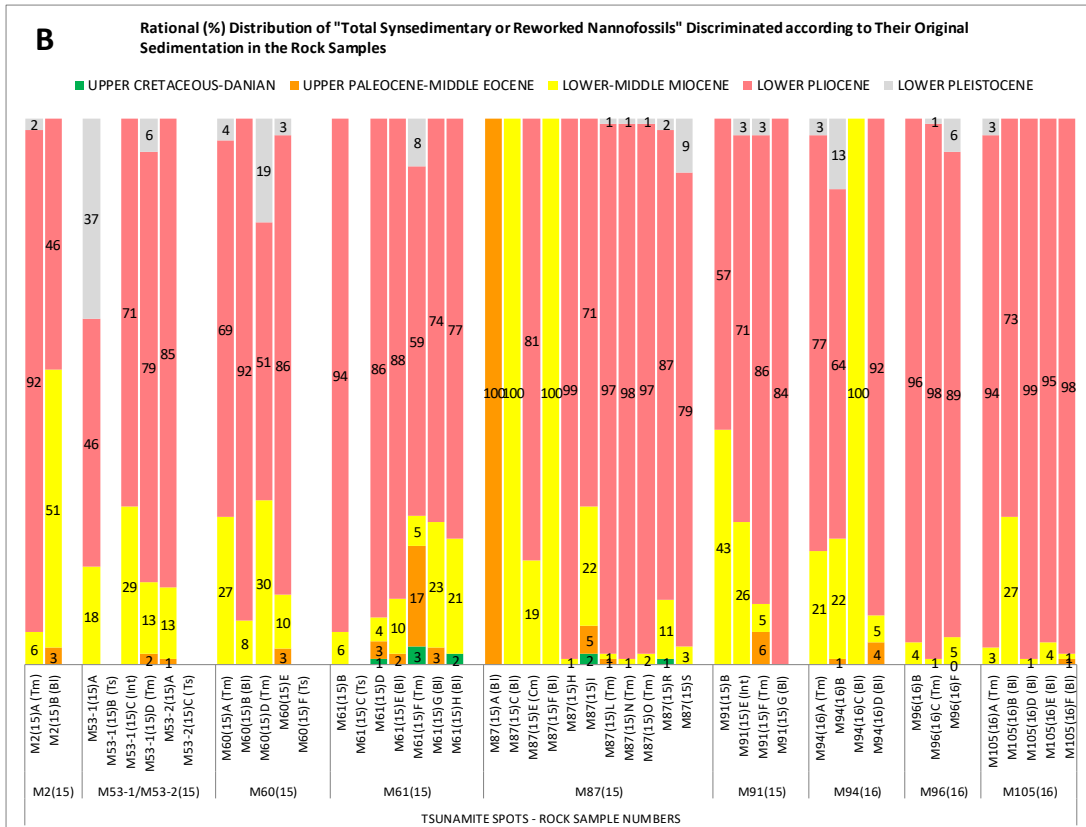
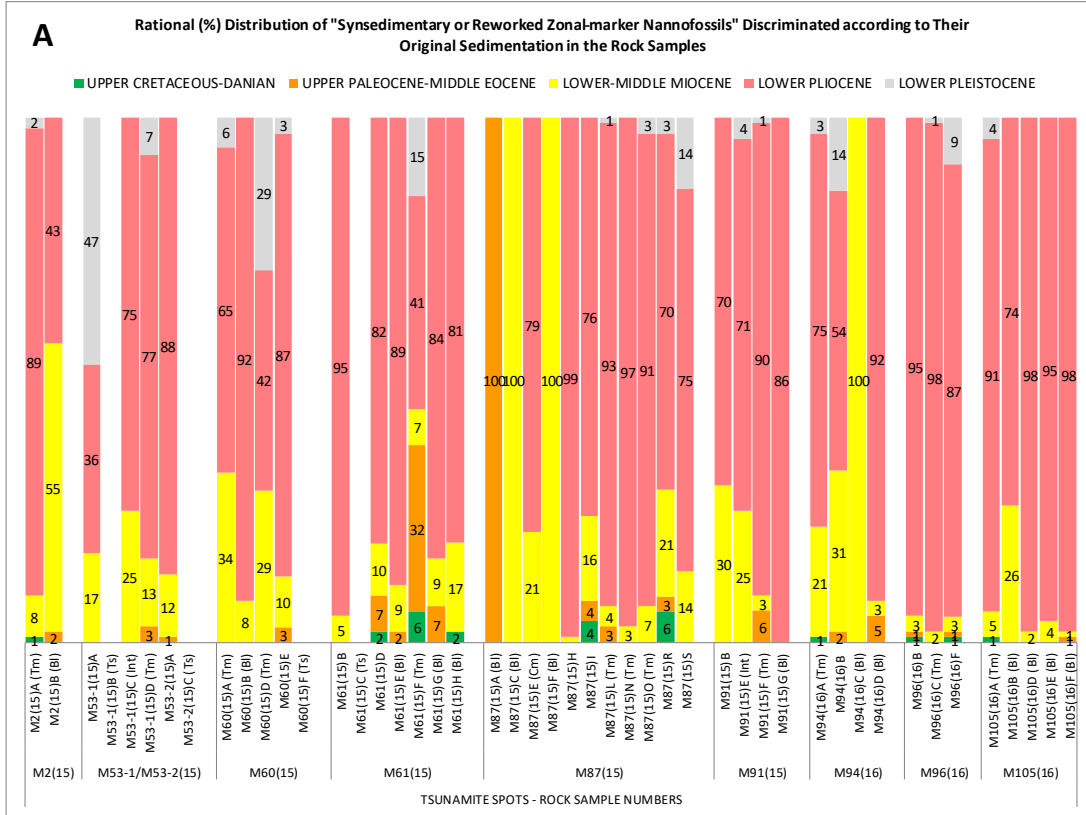


Figure 14. Rational (%) distributions of "synsedimentary or reworked nannofossils" in the Rock Samples of Geceleme, Karpuzçay and Manavgat Formations and tsunami deposits: A) Zonal-marker, B) Total zonal-marker and long-ranged nannofossils discriminated according to their original sedimentations. (fossil quantities are in Tables 1 to 5).

Şekil 14. Geceleme, Karpuzçay ve Manavgat Formasyonları ve tsunami çökellerine ait kayaç örneklerinde "çökelimle eşyaşlı veya yeniden çökeltilmiş nannofosillerin" rasyonel (%) dağılımları: A) Zon-işaretleyici, B) Kökensele çökelimlerine göre ayrılmış toplam zon-işaretleyici ve uzun menzilli nannofosiller. (fosil miktarları Tablo 1 ila 5'te verilmiştir).

Dictyococcites spp. (i.e. *D. hesslandii* (Haq), *D. scripsae* Bukry & Percival), *Pseudotriquetrorhabdulus inversus* (Bukry & Bramlette), *Reticulofenestra* spp. [i.e. *R. dictyoda* (Deflandre in Deflandre & Fert), *R. hillae* Bukry & Percival, *R. reticulata* (Gartner & Smith), *R. stavensis* (Levin & Joerger)], *Sphenolithus moriformis* (Brönnimann & Stradner) are among 22 reworked Middle Eocene nannofossils (see Tab. 4, Fig.15). Upper Cretaceous–Lutetian nannofossils as fine calcareous clastic material repeatedly participated in Early–Middle Miocene, Early Pliocene, and Early Pleistocene sedimentations in addition to synsedimentary and/or other reworked nannofossil data. One Miocene carbonate-clastic succession, Geceleme Formation, supplying reworked nannofossil data into the Early Pliocene and Early Pleistocene sedimentations is located in the study area. Geceleme Formation is represented by NN5 - *Sphenolithus heteromorphus* zone of Langhian–Serravalian ages and have a nannofossil assemblage including *Cyclicargolithus floridanus* (Roth & Hay), *Discoaster druggii* Bramlette & Wilcoxon, *Discoaster minutus* Hojjatzadeh, *Discoaster premicros* de Kaenel & Bergen, *Helicosphaera* spp. [*H. ampliapertura* Bramlette & Wilcoxon, *H. intermedia* Martini, *H. recta* (Haq), *H. stalis* Theodoridis, *H. walbersdorfensis* Müller], *Micrantholithus flos* Deflandre, *Pontosphaera desuetoidea* Bartol, *Reticulofenestra* spp. [*R. haqii* Backman, *R. gelida* (Geitzenauer), *R. minuta* Roth], *Scyphosphaera* spp. [*S. brevis* Varol, *S. lagena* Kamptner, *S. tubifera* Kamptner], *Sphenolithus* spp. [*S. belemnoides* Bramlette and Wilcoxon, *S. compactus* Backman, *S. conicus* Bukry, *S. disbelemnoides* Fornaciari and Rio, *S. dissimilis* Bukry and Percival, *S. heteromorphus* Deflandre], *Triquetrorhabdulus carinatus* Martini, *Umbilicosphaera jafarii* Müller. These 46 synsedimentary nannofossil species with Miocene in addition to the older (reworked from the Cretaceous–Eocene rocks) nannofossils as extrabasinal sedimentary particles are determined to have been reworked



Figure 15. Thematic flow chart explaining sedimentary and stratigraphic origins of different aged nanofossils participating into the rock samples of Geceleme, Karpuzçay and Manavgat Formations and tsunami deposits.

Şekil 15. Geceleme, Karpuzçay ve Manavgat Formasyonları ve tsunami tortullarının kayaç örneklerine katılan farklı yaşlı nanofosillerin sedimentler ve stratigrafik kökenlerini açıklayan konulu akış şeması.

into both Karpuzçay Formation sedimentation (Early Pliocene) and Manavgat Formation (Early Pleistocene) sedimentation (see Tabs. 1–5, Fig.15).

In the rock samples of the Lower Pliocene Karpuzçay Formation were determined 53 synsedimentary nanofossil species including *Calcidiscus* spp. [*C. leptoporus* (Murray & Blackman), *C. macintyreii* (Bukry & Bramlette), *C. tropicus* Kamptner], *Calciosolenia murrayi* Gran, *Coccolithus pliopelagicus* Wise, *Coronosphaera* spp. [*C. binodata* (Kamptner), *C. mediterranea* (Lohmann)], *Discoaster* spp. [*D. asymmetricus* Gartner, *D. brouweri* Tan, *D. pansus* (Bukry & Percival), *D. pentaradiatus* Tan, *D. pliostellulus* Browning & Bergen, *D. surculus* Martini & Bramlette, *D. tamalis* Kamptner, *D. variabilis* Martini & Bramlette], small *Gephyrocapsa* spp., *Helicosphaera* spp. [*H. carteri* (Wallich), *H. kamptneri* Hay & Mohler, *H. wallichii* (Lohmann), *H. selli* (Bukry and Bramlette)], *Pontosphaera* spp. [*P. discopora* Schiller, *P. multipora* (Kamptner), *P. japonica* (Takayama)], *Pseudoemiliana lacunosa ovata* (Kamptner), *Reticulofenestra* spp. [*R. minutula* (Gartner), *R. pseudoumbilica* (Gartner)], *Rhabdosphaera clavigera* Murray & Blackman, *Scyphosphaera* spp. [*S. apsteinii* Lohmann, *S. canescens* Kamptner, *S. recurvata* Deflandre, *S. pulcherrima* Deflandre], *Sphenolithus* spp. [*S. abies* Deflandre, *S. U. neoabies* Bukry & Bramlette], *Sphenaster metula* (Wilcoxon), *Syracosphaera pulchra* Lohmann, *Umbilicosphaera* spp. [*U. foliosa* Kamptner, *U. rotula* Kamptner, *U. sibogae* (Weber-van Bosse)] and 3 dinoflagellate cysts are found *Calciodinellum albatrosianum* (Kamptner), *Lebessphaera urania* Meier, Janofske & Willems, *Pernambugia tuberosa* (Kamptner).

All the Pliocene nanofossil assemblage together with the older aged nanofossil data indicating to the Upper Cretaceous-Middle Miocene marine sedimentations are also seen as an abundant reworked nanofossil data participating into the Lower Pleistocene aged Manavgat Formation deposits (see Tabs. 1–5, Fig.15). These data represent NN15 (Martini, 1971) or CN11 (Okada and Bukry, 1980) *Reticulofenestra pseudoumbilicus* zone of Late Zanclean between ~3,98 and ~3,45 ma. This zone corresponds to CNPL3 – *Discoaster asymmetricus* / *Reticulofenestra pseudoumbilicus* concurrent range zone according to Backman et al. (2012) and CNPL4 *Discoaster tamalis* top zone transition (Fig. 16).

Tsunami deposits and regular sedimentary rocks of the Manavgat Formation include an abundant reworked nanofossil data eroded from the older marine sedimentary rocks of

four stratigraphic sequences placed between Upper Cretaceous-Lower Pliocene in addition to the symsedimentary nannofossils such as *Calcidiscus leptoporus* (Murray & Blackman), *Ceratolithus cristatus* HET coccolithomorpha (Lecal-Schlauder), middle *Gephyrocapsa* spp. [*G. caribbeana* Boudreaux & Hay, *G. lumina* Bukry], small *Gephyrocapsa* spp., *Helicosphaera* spp. [*H. hyalina* Gaarder, *H. pavementum* Okada & McIntyre], *Oolithotus fragilis* (Lohmann), *Pontosphaera scutellum* Kamptner, *Pseudoemiliana* spp. [*P. lacunosa* (Kamptner), *P. ovata* (Kamptner)], *Scyphosphaera* spp. [*S. expansa* Bukry & Percival, *S. porosa* Kamptner, *Umbellosphaera tenuis* (Kamptner), *Umbilicosphaera* spp. [*U. foliosa* Kamptner, *U. sibogae* (Weber-van Bosse)], *Leonella granifera* (Fütterer) (Tabs. 1–5, Fig.15). In addition to these Lower Pleistocene nannofossil assemblage within the tsunamite matrix, an abundant Lower Pliocene, and Lower-Middle Miocene, Upper Paleocene–Middle Eocene nannofossil data involve both in the matrix and in the blue marl pebble/cobble/blocks of Karpuzçay, Geceleme Formations or the older sequences.

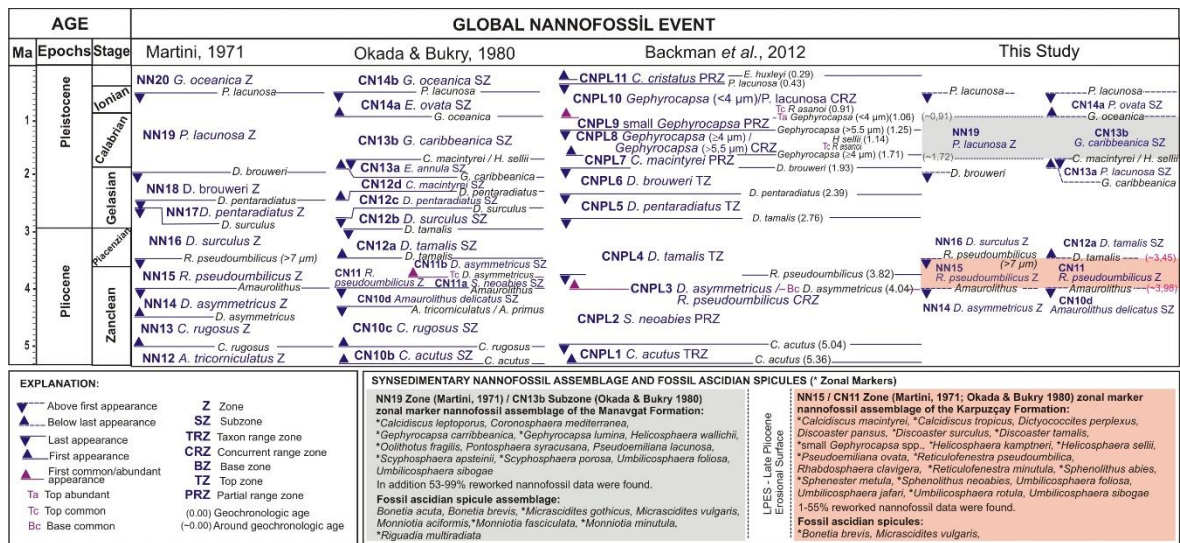


Figure 16. In this study, compared with Martini (1971), Okada and Bukry (1980) and Backman *et al.* (2012), Early Pliocene and Early Pleistocene nanoplankton zones and chrono-bio events of Karpuzçay and Manavgat (including Tusanami sediments) Formations.

Şekil 16. Bu çalışmada, Martini (1971), Okada ve Bukry (1980) ve Backman ve diğerleri (2012) ile karşılaştırmalı olarak, Karpuzçay ve Manavgat (Tusanami çökelleri dahil) Formasyonlarının Erken Pliyosen ve Erken Pleistosen nanoplankton zonları ve krono-biyostratigrafik olguları.

Although the relative presence of Pliocene nannofossil assemblage as reworked sedimentary material including *Calcidiscus macintyre* (Bukry & Bramlette), *Discoaster*

spp. [i.e. *D. asymmetricus* Gartner, *D. brouweri* Tan, *D. pansus* (Bukry & Percival), *D. pentaradiatus* Tan, *D. pliostellulus* Browning & Bergen, *D. surculus* Martini & Bramlette, *D. tamalis* Kamptner], the most of small *Gephyrocapsa* spp., *Helicosphaera* spp. [i.e. *H. carteri* (Wallich), *H. kamptneri* Hay & Mohler, *H. wallichii* (Lohmann), *H. selli* (Bukry and Bramlette)], *Reticulofenestra pseudoumbilica* (Gartner), *Scyphosphaera* spp. [*S. apsteinii* Lohmann, *S. canescens* Kamptner, *S. intermedia* Deflandre, *S. recurvata* Deflandre, *S. pulcherrima* Deflandre], *Sphenolithus* spp. [*S. abies* Deflandre, *S. neoabies* Bukry & Bramlette], *Sphenaster metula* (Wilcoxon), *Umbilicosphaera* spp. [*U. foliosa* Kamptner, *U. jafarii* Müller, *U. rotula* Kamptner, *U. sibogae* (Weber-van Bosse)], determined sysedimentary nannofossils such as *Calcidiscus leptoporus* (Murray & Blackman), *Ceratolithus cristatus* HET coccolithomorpha (Lecal-Schlauder), *Gephyrocapsa* spp. [i.e. *G. caribbeanica* Boudreaux & Hay, *G. lumina* Bukry], *Helicosphaera* spp. [i.e. *H. hyalina* Gaarder, *H. pavementum* Okada & McIntyre], *Oolithotus fragilis* (Lohmann), *Scyphosphaera* spp. [i.e. *S. apsteinii* Lohmann, *S. expansa* Bukry & Percival, *S. porosa* Kamptner], *Umbellosphaera tenuis* (Kamptner) correspond into NN19 - *Pseudoemiliana lacunosa* Zone (Martini, 1971) or CN13b - *Gephyrocapsa caribbeanica* Subzone (Okada and Bukry, 1980) of Early Pleistocene (Calabrian) between ~1,72 and ~0,91 ma. This zone does not correspond to CNPL8 – *Gephyrocapsa* ($\geq 4 \mu\text{m}$) / *Gephyrocapsa* ($> 5.5 \mu\text{m}$) Concurrent Range Zone according to Backman et al. (2012) (see Tabs. 1–5, Fig. 16).

6. DISCUSSION AND RESULTS

The Mediterranean region is seismically active due to the convergence of the African Plate with the Eurasian plate and Present day Africa-Eurasia motion ranges from ~4 millimetres per year (mm/yr) in a northwest-southeast direction in the western Mediterranean to ~10 mm/yr (north-south) in the eastern Mediterranean according to Herman et al. (2015). They have stated that the Africa-Eurasia convergence began approximately 50 million years ago and was associated with the closure of the Tethys Sea; the Mediterranean Sea is all that remains of the Tethys. The highest rates of seismicity in the Mediterranean region were described along the Hellenic subduction zone of southern Greece and significant rates of current seismicity and at the southern edge of the Anatolian Block lies the east-west trending Cyprian Arc, which hosts moderate levels of seismicity.

The Cyprian Arc represents the convergent boundary between the Anatolian Block to the north and the Africa Plate to the south and large historical earthquakes have occurred throughout the region spanning the Mediterranean Sea. As being similar to today-cases, a tsunami event must have occurred most likely due to emerging a new earthquake fracturing at the bottom of the Mediterranean Sea in the early Pleistocene due to the Africa-Eurasia convergence. Consequently, a new marine sedimentation which built up shallow marine clastics of the Manavgat Formation was occurred as a result of the Lower Pleistocene tsunami event continuing as a sea-surfing up and occupying over the Late Pliocene Erosional Surface (LPES) on the Manavgat coastal plane. In former studies, the originally indiscriminated nannofossil data directly used for chronostratigraphy caused complexity related to stratigraphic description between Neogene and Quaternary sequences. It is concluded that Karpuzçay Formation, which was previously, determined as Burdigalian–Tortonian aged, in fact, was divided into Lower Pliocene and Lower Pleistocene aged two formations belonging to two different marine sedimentations.

Tsunami deposits “Manavgat Formation Tsunami Member” include a complex nannofossil assemblage supplied from four different original sources: (1) synsedimentary data deposited inner shelf, (2) synsedimentary data upwelling from deep sea bottom, (3) removed data derived from the intrabasin, (4) reworked data derived from the extrabasin. Synsedimentary data consist of Lower Pleistocene nannofossils representing NN19 Zone (Martini, 1971)/CN13b Subzone (Okada and Bukry, 1980) in the pebbly mud matrix or rip-up mud clasts of the tsunamite. As for reworked data, Upper Cretaceous-Lower Pliocene aged nannofossil data should be accepted as stratigraphically misleading old sedimentary particules either eroded from the different sedimentary rocks or coming from the muddy bottom sediments.

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8. REFERENCES

- Akay, E. and Uysal, S., 1985, Stratigraphy, sedimentology and structural geology of Neogene sediments at the western of the Central Taurides (Antalya). Mineral Research and Exploration Institute, Report No. 7799, 151 p. (in Turkish, unpublished)
- Akay, E., Uysal, S., Poisson, A., Cravatte, J. and Müller, C., 1985, Stratigraphy of the Antalya Neogene Basin. Geological Society Bulletin of Turkey, 28,105–119. (in Turkish with English abstract)
- Akbulut, A., 1977, Etüde géologique d'une partie du Taurusoccidental au Sud d'Eğridir (Turquie): These 3^oCycle Université Paris-Sud Orsay, 203 p.
- Altınlı, E., 1943, Geology of the Antalya Region. Mineral Research and Exploration Institute, Report No. 6858–6859, 205 p. (in Turkish, unpublished)
- Backman, J., Raffi, I., Rio, D., Fornaciari, E. and Pälke, H. 2012, Biozonation andbiochronology of Miocene through Pleistocene calcareous nannofossils from low and middle latitudes. Newsletters on Stratigraphy, 45, 221–244.
<http://doi.org/10.1127/0078-0421/2012/0022>
- Bown, P.R. and Young, J.R., 1998, Techniques, in: Bown, P.R. (Ed.), Calcareous Nannofossil Biostratigraphy. British Micropalaeontological Society Publications Series. Chapman and Hall, London, ISBN: 0412789701, 16–28.
- Çiner, A., Karabıyıköğlü, M., Monod, O., Deynoux, M. and Tuzcu, S., 2008, Late Cenozoic sedimentary evolution of the Antalya Basin, Southern Turkey. Turkish Journal of Earth Sciences, 17, 1–41. <https://doi.org/10.4154/gc.2022.30>
- Dawson, A.G. and Stewart, I., 2007, Tsunami deposits in the geological record. Sedimentary Geology, 200, 166–183.
<https://doi.org/10.1016/j.sedgeo.2007.01.002>
- Deynoux, M., Çiner, A., Monod, O., Karabıyıköğlü, M., Manatschal, G. and Tuzcu, S., 2005, Facies architecture and depositional evolution of alluvial fan to fan delta complexes in the tectonically active Miocene Köprü çay basin, Isparta angle, Turkey. In: Kelling, G., Robertson, A.H.F. and Buchem, F. (Eds), Cenozoic

- sedimentary basins of South Central Turkey. *Sedimentary Geology*, 173, 315–343. <https://doi.org/10.1016/j.sedgeo.2003.12.013>
- Eroskay, O., 1968, Geological investigation of the Köprüçay-Beşkonak Reservoir area, Electrical Works and Etudes Administration, Ankara, Report No. II-06-5, 80 p.
- Flecker, R., Ellam, R.M., Müller, C., Poisson, A., Robertson, A.H.F. and Turner, J., 1998, Application of Sr isotope stratigraphy and sedimentary analysis to the origin and evolution of the Neogene Basins in the Isparta Angle, Southern Turkey. *Tectonophysics*, 298, 83–101. [https://doi.org/10.1016/S0040-1951\(98\)00179-6](https://doi.org/10.1016/S0040-1951(98)00179-6)
- Flecker, R., Poisson, A. and Robertson, A.H.F., 2005, Facies and palaeogeographical evidence for the Miocene evolution of the Isparta Angle in its regional Eastern Mediterranean context. *Sedimentary Geology*, 173, 277-314. <https://doi.org/10.1016/j.sedgeo.2003.10.014>
- Fujiwara O. and Kamataki, T., 2007, Identification of tsunami deposits considering the tsunami waveform: An example of subaqueous tsunami deposits in Holocene shallow bay on southern Boso Peninsula, Central Japan. *Sedimentary Geology*, 200, 295–313.
- Glover, C. and Robertson, A.H.F., 1998a, Role of regional extensional and uplift in the Plio-Pleistocene evolution of the Aksu Basin, SW Turkey. *Journal of the Geological Society*, London, 15, 365–388. <https://doi.org/10.1144/gsjgs.155.2.0365>
- Glover, C. and Robertson, A.H.F., 1998b, Neogene intersection of the Aegean and Cyprus arcs: extensional and strike-slip faulting in the Isparta Angle, SW Turkey. *Tectonophysics*, 298, 103–132. [https://doi.org/10.1016/S0040-1951\(98\)00180-2](https://doi.org/10.1016/S0040-1951(98)00180-2)
- Görmüş, M., Sagular, E.K. and Çoban H., 2001, The Miocene sequence characteristics, its contact relation to the older rocks and lamprophyric dikes in the Dereboğazı area (S Isparta, Turkey). 4th. International Symposium on Eastern Mediterranean Geology, Isparta, Turkey, 21–25 May 2001, Proceedings, p. 69–90.

- Gutnic, M., Monod, O., Poisson, A. and Dumont, J.F., 1979, Géologie des Taurides Occidentales (Turquie). Mémoires de la Société géologique de France, 58(137), 1–112.
- Hepdeniz, K. and Sagular, E.K., 2009, Stratigraphical Analysis Based Upon Generic Discrimination of the Nannofossil Data of the Tertiary-Aged Marine Clastics Around the District of Direkli and Darıdere Villages (South of Isparta). Süleyman Demirel University Journal of Natural and Applied Sciences, 13(3), 245–264. (in Turkish with English abstract)
- Herman, M.W., Hayes, G.P., Smoczyk, G.M., Turner, R., Turner, B., Jenkins, J., Davies, S., Parker, A., Sinclair, A., Benz, H.M., Furlong, K.P., and Villaseñor, A., 2015, Seismicity of the Earth 1900–2013, Mediterranean Sea and vicinity: U.S. Geological Survey Open File Report 2010–1083-Q, scale 1:10,000,000. <https://doi.org/10.3133/ofr20101083Q>
- Karabıyıkoglu, M.F., Çiner A., Monod, O., Deynoux, M., Tuzcu, S. and Örcen, S., 2000, Tectono-sedimentary evolution of the Miocene Manavgat Basin, western Taurides, Turkey, in Bozkurt, E., Winchester, J.A., Piper, J.D.A. (Eds), Tectonics and Magmatism in Turkey and the Surrounding Area, Geological Society, London, Special Publication, 173, 271–295. <https://doi.org/10.1144/GSL.SP.2000.173.01.14>
- Karabıyıkoglu, M., Tuzcu, S., Çiner, A., Deynoux, M., Örcen, S. and Hakyemez, A., 2005, Facies and environmental setting of the Miocene coral reefs in the late-orogenic fill of the Antalya Basin, Western Taurides, Turkey: Implications for the Tectonic Control and sea-level changes. *Sedimentary Geology*, 173, 345–371. <https://doi.org/10.1016/j.sedgeo.2003.08.006>
- Kortekaas, S. and Dawson, A.G., 2007, Distinguishing tsunami and storm deposits: An example from Martinhal, SW Portugal. *Sedimentary Geology*, 200, 295–313. <https://doi.org/10.1016/j.sedgeo.2007.01.004>
- Martini, E., 1971, Standard Tertiary and Quaternary calcareous nannoplankton zonation, in Proceedings of the Second Planktonic Conference Roma 1970, (ed. A. Farinacci), Edizioni Technoscienza, Rome, 2, 739–785.

- Monod, O., 1977, Recherches Geologiques dans le Taurus occidental au Sud de Beyşehir (Turquie): These d'Etat Univ. Paris Sud (Orsay), 442 p.
- Morton, R.A., Gelfenbaum, G. and Jaffe, B.E., 2007, Physical criteria for distinguishing sandy tsunami and storm deposits using modern examples. *Sedimentary Geology*, 200, 184–207. <https://doi.org/10.1016/j.sedgeo.2007.01.003>
- Okada, H. and Bukry, D., 1980, Supplementary modification and introduction of code numbers to the low-latitude coccolith biostratigraphic zonation (Bukry, 1973:1975). *Marine Micropaleontology*, 5 (3), 321–325. [https://doi.org/10.1016/0377-8398\(80\)90016-X](https://doi.org/10.1016/0377-8398(80)90016-X)
- Özbakır, A.D., Şengör, A.M.C., Wortel, M.J.R., and Govers, R., 2013, The Pliny–Strabo trench region—A large shear zone resulting from slab tearing. *Earth and Planetary Science Letters*, 375, 188–195. <http://doi.org/10.1016/j.epsl.2013.05.025>
- Perch-Nielsen, K. 1985a, Mesozoic calcareous nannofossils. Bolli, H. M., Saunders, J. B., Perch-Nielsen, K. (Ed.). *Plankton Stratigraphy*, Cambridge Earth Sciences Series, p. 329–426.
- Perch-Nielsen, K. 1985b, Cenozoic calcareous nannofossils. Bolli, H. M., Saunders, J. B., Perch-Nielsen, K. (Ed.). *Plankton Stratigraphy*, Cambridge Earth Sciences Series, p. 427–554.
- Poisson, A., 1977, Recherches geologiques dans les Taurides occidentales (Turquie): These d'Etat Univ.Paris-Sud (Orsay), 795 p. (unpublished).
- Poisson, A., Wernli, R., Sagular, E.K. and Temiz, H., 2003b, New data concerning the age of the Aksu Thrust in the south of the Aksu valley, Isparta Angle (SW Turkey): consequences for the Antalya Basin and the Eastern Mediterranean. *Geological Journal*, 38, 311–327. <https://doi.org/10.1002/gj.958>
- Robertson, A.H.F., 1998, Mesozoic–Tertiary tectonic evolution of the Easternmost Mediterranean area: integration of marine and land-based evidence, in Robertson A.H.F., Emeis K.C., Richter C., Camerlenghi A. (eds), *Proceedings of the Ocean Drilling Program, Scientific Results*, 160, 723–782. (available at http://www-odp.tamu.edu/publications/160_SR/160TOC.HTM/)

- Robertson, A.H.F., Poisson, A. and Akıncı, A., 2003, Developments in research concerning Mesozoic-Tertiary Tethys and Neotectonics in the Isparta Angle, SW Turkey. *Geological Journal*, 38 (3–4), 195–234. <https://doi.org/10.1002/gj.953>
- Sagular, E.K., 2003. A new determination method on using nannofossil data in stratigraphic ageing and paleoenvironmental analyses. *Suleyman Demirel University Journal of Natural and Applied Sciences, Special Issue*, 7 (2), 25–36. (in Turkish with English abstract)
- Sagular, E.K., 2009, Fossil didemnid ascidian spicule records in the Plio-Quaternary marine clastics of the Antalya basin (Eastern Mediterranean) and their stratigraphic calibration to new nannofossil data. *Geosciences Journal*, 13 (2), 121–131. <http://doi.org/10.1007/s12303-009-0011-7>
- Sagular, E.K., Yavuzlar, G., Aydemir, E. and Yüzgöl, N.S., 2016, Stratigraphic-sedimentologic significance of genetically analysed nannofossil records in sedimentary rocks: Examples from the Middle Miocene - Early Pleistocene marine deposits of Manavgat Basin (SW Turkey). *The Micropaleontological Society, TMS Foraminifera and Nannofossil Groups Joint Meeting 2016*, 19-24 June 2016, Angers, France, Meeting Handbook and Abstracts, p. 160.
- Sagular, E.K. and Yavuzlar, G., 2017, Calcareous nannofossils from the Manavgat Sub-basin, SW Turkey, reveal the age of lower Pleistocene paleotsunami deposits that overlie a late Pliocene erosional surface. *International Nannoplankton Association Meeting*, 24-28 September 2017, Athens, Greece, *JNR 37 September 2017* (special issue), Abstracts, p. 143.
- Sagular, E.K., Yümün, Z.Ü. and Meriç, E., 2018, New didemnid ascidian spicule records calibrated to the nannofossil data chronostratigraphically in the Quaternary marine deposits of Lake Iznik (NW Turkey) and their paleoenvironmental interpretations, *Quaternary International*, 486, 143–155. <http://dx.doi.org/10.1016/j.quaint.2017.08.060>
- Şenel, M., 1997, 1:100000 scaled geological maps of Turkey: Antalya-L12 (O26) section. *General Directorate of Mineral Research and Exploration, Publication No. 9*, 14 p., (in Turkish with English abstract).

- Tarr, A.C., Villaseñor, A., Furlong, K.P., Rhea, S., and Benz, H.M., 2010, Seismicity of the Earth 1900–2007: U.S. Geological Survey Scientific Investigations Map 3064, 1 sheet, scale 1:25,000,000. <https://doi.org/10.3133/sim3064>
- Varol, O. and Houghton, S.D., 1996. A review and classification of fossil didemnid ascidian spicules. *Journal of Micropalaeontology*, 15, 135–149. <https://doi.org/10.1144/jm.15.2.1>
- Yavuzlar, G. and Sagular, E.K., 2018, Revised Chrono-biostratigraphy of Lower Miocene deposits of the Eastern Mediterranean (SW Turkey), based on calcareous nanofossils. *Geologia Croatica*, 71(3), 121–133. <https://doi.org/10.4154/gc.2018.15>
- Young, J.R., Bown P.R. and Lees J.A., 2017, Nannotax3 website. International Nannoplankton Association. <http://www.mikrotax.org/Nannotax3> (accessed 21 Apr. 2017).
- https://earthquake.usgs.gov/earthquakes/tectonic/images/mediterranean_tsum.pdf