



# Bulletin of the Mineral Research and Exploration

<http://bulletin.mta.gov.tr>



## First Turolian findings in the Neogene sequence of Denizli Basin (SW Anatolia) and its regional palaeobiogeographic significance

Adil DOĞAN<sup>a\*</sup>, Serdar MAYDA<sup>b</sup> and M. Cihat ALÇİÇEK<sup>c</sup>

<sup>a</sup>General Directorate Mineral Research and Exploration, 2<sup>nd</sup> Central Anatolian Regional Directorate, Konya, Turkey

<sup>b</sup>Ege University, Department of Biology, İzmir, Turkey

<sup>c</sup>Pamukkale University, Department of Geology, Denizli, Turkey

Research Article

### Keywords:

Terrestrial Neogene,  
Vertebrate, Turolian,  
Denizli, Western Anatolia.

### ABSTRACT

The large fossil vertebrates obtained from the alluvial flood-plain deposits of the Kolankaya formation are determined as *Skoufotragus laticeps* (Andree, 1926) and *Hipparion brachypus* (Hensel, 1862), as representative elements of palaeomammal faunas spanning from the eastern Mediterranean to Iranian domains during the late Miocene (early-middle Turolian, MN11-12). These first Turolian records from the basin fill succession bear importance on reconstruction for the palaeobiogeographic diversity of relevant taxons as well as admit of interbasinal stratigraphic correlation for the Western Anatolian terrestrial Neogene basins.

Received Date: 20.07.2019

Accepted Date: 25.11.2019

## 1. Introduction

The Anatolian peninsula has a key position for migration and distribution of vertebrate groups due to its geographic location between Asia, Africa and Europe. Therefore, the rich fossil content of the terrestrial Neogene basins of Anatolia is of great importance in understanding inter-continental faunal changes and biodiversity (Sickenber and Tobien, 1971; Saraç, 2003; Alçiçek, 2010; Alçiçek et al., 2019; Krijsgman et al., 2019). Although there are already known vertebrate fossil records in the Neogene sequence of the Denizli Basin (Saraç, 2003), yet there is no record of Turolian time. On the other hand, the Turolian vertebrate fossil records are already known in many localities around the Denizli Basin such as Asarlık (Alçiçek et al., 2012, 2019; Alçiçek and Alçiçek, 2014), Mahmutgazi (Sickenberg et al., 1975; Pickford and Ertürk, 1979; Köhler, 1987; Saraç, 2003; Pickford, 2016; Geraads, 2017), Sazak (Saraç, 2003; Tuna, 1999), Elmalıyurt

(Saraç, 2003; Alçiçek et al., 2005, 2006, 2019); Kemiklitepe (Şen et al., 1994; Saraç, 2003; Xafis et al., 2020), Karabeyli (Seyitoğlu et al., 2009), Karamusalar (Alçiçek, 2007; Alçiçek et al., 2019), Şerefköy (Saraç, 2003; Kaya et al., 2012; Mayda, 2014), Özlüce (Saraç, 2003; Alpagut et al., 2014; Mayda, 2014).

In this study, the presence of regional inter-basinal faunal connectivity and migration routes have been shown with support of the first Turolian vertebrates found in the Denizli Neogene sequence for the first time. These findings, which fill an important gap for a more detailed understanding of Turolian palaeobiogeography in Anatolia, also enable regional stratigraphic correlation among the basins.

## 2. Geological Setting

Since the Miocene, the western Anatolia has been consisted of horst-graben system driven by the

Citation info: Doğan, A., Mayda, S., Alçiçek, M. C. 2020. First Turolian findings in the Neogene sequence of Denizli Basin (SW Anatolia) and its regional palaeobiogeographic significance. Bulletin of the Mineral Research and Exploration 163, 1-12.  
<https://doi.org/10.19111/bulletinofmre.651620>.

\*Corresponding author: Adil DOĞAN, [adil.dogan@mta.gov.tr](mailto:adil.dogan@mta.gov.tr)

regional crustal extension and associated normal faults. The metamorphics of the Precambrian-Paleozoic Menderes Massif and tectonically overlying the Mesozoic carbonates and ophiolites of the Lycian Nappes are the basement rock suits. The regional extension was resulted in the subsidence of the Denizli Basin from the beginning of the early Miocene and initiated alluvial, fluvial and lacustrine sedimentation (Şimşek, 1984; Okay, 1989; Sun, 1990; Westaway, 1993; Konak and Şenel, 2002; Koçyiğit, 2005; Kaymakçı, 2006; Alçiçek et al., 2007; figure 1). Westaway (1993) states the Denizli Basin as a graben that subsided continually by the extension of western Anatolia, while Koçyiğit (2005) proposes a short term of compression interrupted the extension. Based on micro-mammal fossils Saraç (2003) attributes an early Miocene age to the oldest basin-fill deposits.

Kaymakçı (2006) states that the extension in the basin was accompanied by transfer faults and documented a late Pliocene age of the basin based on micro-mammal fossils. According to Alçiçek et al. (2007) sedimentation in the basin has been constant until present in control of NW trending faults that limit the basin and accompanied by the climatic and lake-level changes.

In this study, a detailed geological map of the study area to the northeast of Denizli Basin framed in the Denizli M22-a2 sheet has been produced in 1/25.000 scale (Figure 2). The stratigraphy of the basin was elaborated by correlating the findings of the new Turolian fauna in the study area with the previously founded fossils in surrounding basins. The new fossils were found at Gölyeri (UTM 693655/4200295) within the Güzelpınar member of Kolankaya formation which consists of alluvial-fan deposits, tectonically elevated at the northeast margin of the basin.

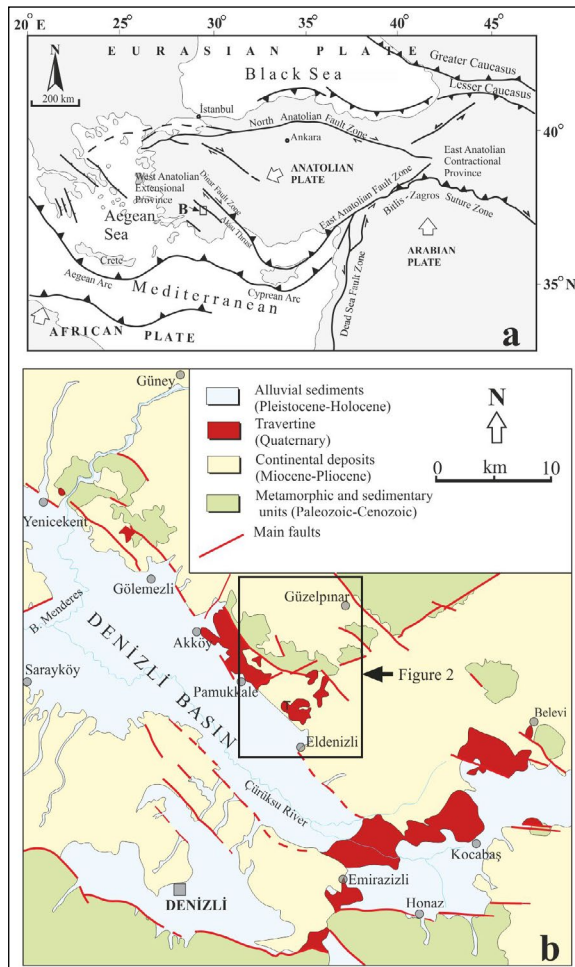


Figure 1- a) Regional tectonic setting of Anatolia (Bozkurt, 2003; Alçiçek et al., 2013; Kaymakçı et al., 2018) and b) geological map of the Denizli Basin (Sun, 1990).

### 3. Neogene Stratigraphy of the Denizli Basin

The Neogene units that unconformably overlie the basement in the study area were first described by Şimşek (1984) as the Denizli Group (Figure 3).

#### 3.1. Kızılburun Formation

The unit forming the base of Denizli Neogene succession overlies unconformably the basement rocks. It is consisted of conglomerate at the base and sandstone, siltstone and mudstone alternations towards the top. The upper levels include clayey limestones, siltstones, mudstones and coal lenses. The unit is conformably overlain by lacustrine Sazak formation. The age of the unit was determined as early Miocene by Saraç (2003) according to the micro-mammalian remains (MN5-6).

#### 3.2. Sazak Formation

The unit is mainly composed of lacustrine sediments, consists of limestones, marls, siltstones, claystones and organic mudstones, cherty limestones which contain gypsum, gypsum arenite and organic levels. The unit conformably overlies the Kızılburun formation as well as onlaps the basement units, and overlies by the Kolankaya formation. Saraç (2003) suggests a middle Miocene age for the unit according to the micro-mammalian associations (MN6-8).

### 3.3. Kolankaya Formation

The dominant lithology of the unit is yellowish-brown, gray coloured sandstones with altered sandstones, marls and limestones at the base. Thin-medium to thick bedded sandstones are loosely

cemented, dispersed and bright mica flakes within them are among the characteristics features in the field. Interlayers of clayey limestone are common in the unit and gypsum layers developed in the marshy environment towards the top. Fluvial deposits reaching

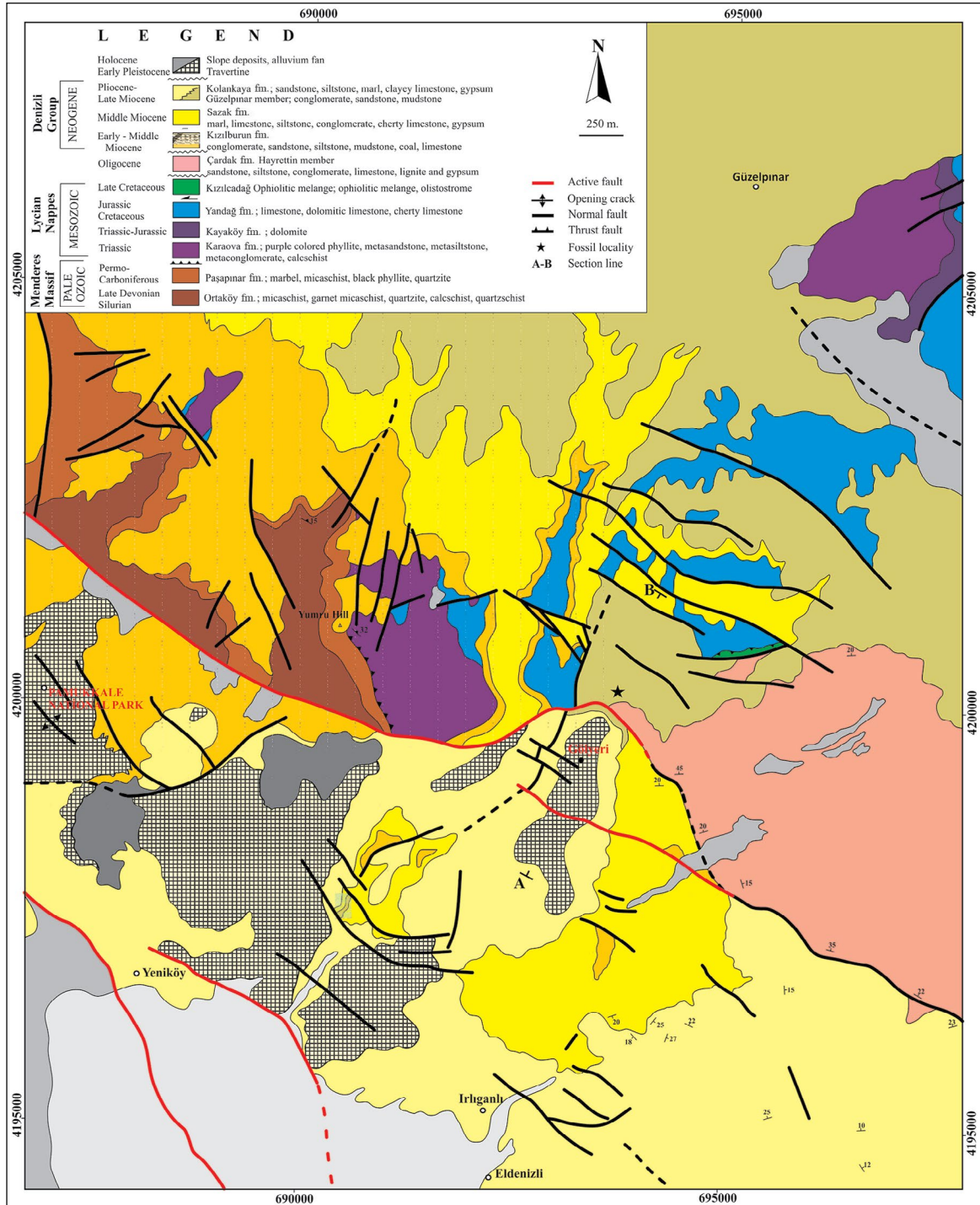


Figure 2- Detailed geological map of the study area (Denizli M22-a2). Based on the unpublished geological map of MTA in 1/25.000 scale compiled by N. Konak, and the active tectonic map of MTA (Emre et al., 2013).

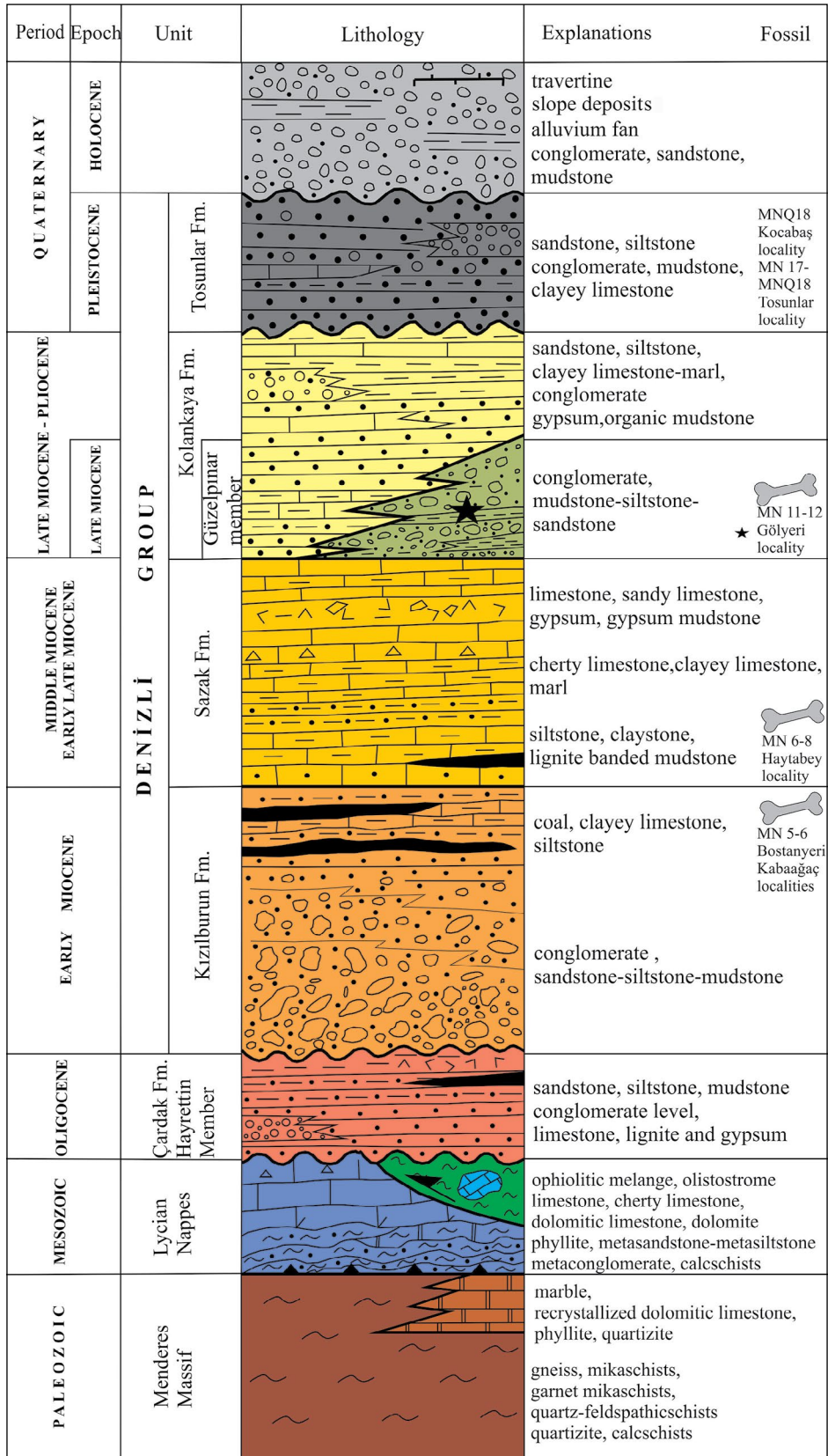


Figure 3- Stratigraphy of the Denizli Basin (Nebert, 1958; Şimşek, 1984; Okay, 1989; Sun, 1990; Alçiçek et al., 2007; Kaymakçı, 2006; Saraç, 2003; Lebatard et al., 2014a, b; Boulbes et al., 2014; Alçiçek et al., 2015; Rausch et al., 2019).

into the lacustrine part of the unit from the basin margins form shallow water deltaic successions. Saraç (2003) attributes a late Miocene-late Pliocene age to the unit based on the micro-mammal associations.

### 3.3.1. Güzelpınar Member

The fossil bearing unit presented in this study is defined as Güzelpınar member within the Kolankaya formation. The unit dominated by red, reddish brown coloured irregular alternation of conglomerate, sandstone, siltstone and mudstone. The pebbles are polygenic and rounded; angular to semi-angular pebbles are also observed. Sandstone-siltstone levels commonly contain caliches. The poorly-sorted, mass-

transport dominated unit is mainly composed of alluvial flood deposits (Figure 4). The age of the unit is assigned to late Miocene (early-middle Turolian, MN11-12) according to the vertebrate fossils introduced in this study.

### 3.4. Tosunlar Formation

The dominant lithology of the unit is composed of yellowish-brownish sandstone and siltstones, conglomerate, mudstone and clayey limestones. It constitutes the uppermost reach of the Neogene sequence of the basin. The age of the unit has been determined as late Pleistocene based on the micro-mammal content of (MN17) by Kaymakçı (2006).

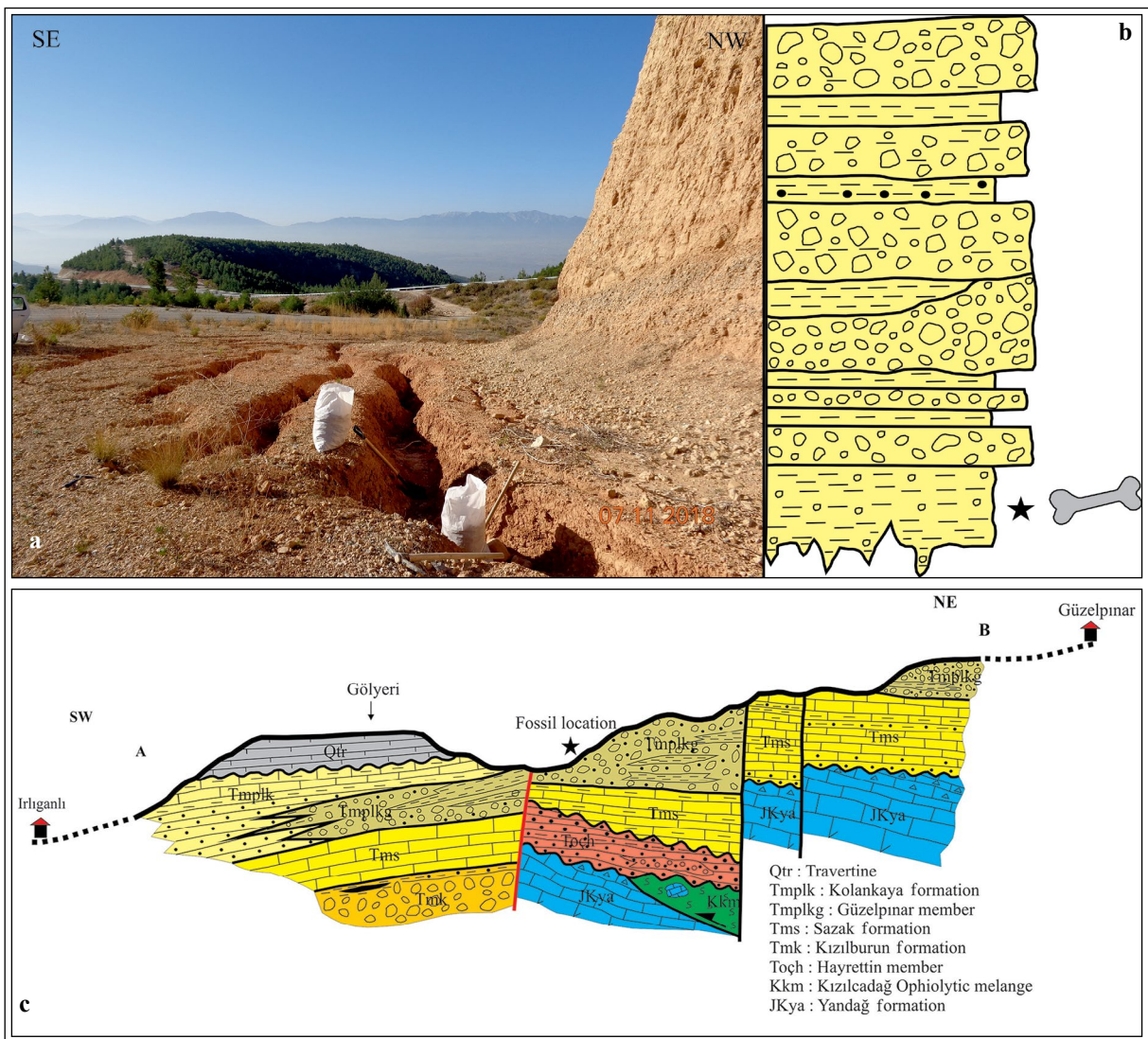


Figure 4- a) Field view of the Turolian fossil locality at Gölyeri, b) stratigraphic section of the fossil bearing unit and c) geological cross section through the fossil locality (not to scale).

The youngest units in the study area are the undifferentiated Quaternary alluvial-fan and scree and travertine deposits.

#### 4. Material and Method

The fossil materials in this study (Figure 5) were found in situ and recorded in the inventory of Ege University Natural History Museum (EUNHM). The odontological study is based on a description of the structural elements of the tooth and comparison with the definitions in the previous studies. In this respect, for the dental terminology of the Bovidae, Gentry et al. (1999) has been followed. The upper teeth are indicated with M (upper case) and the measurements are obtained by measuring the maximum length and perpendicular width with a 0.1 mm precision calliper. All measurements are given in millimeters (mm). For the method and terminology of *Hipparion* metacarpal measurement, Eisenmann et al. (1988) has been followed. Measurements: M10 (maximum distal supra-articular diameter), M11 (maximum distal articular diameter), M12 (maximum diameter of the distal keel), M13 (minimum diameter of the distal lateral conduit), M14 (maximum diameter of the distal medial conduit) have been calculated and the values are depicted graphically in figure 6 and figure 7. Şen et al. (1978) suggested that the sagittal keel

development observed in the distal third metapodials of Equidae family is an important morphological character to distinguish between the Vallesian and Turolian *Hipparion* forms, and they first revealed the index which was defined as the keel index. This index is obtained by (M12x100/M13) operation and this value increases from the Vallesian to Turolian. The development of the sagittal keel is directly related to the lateral mobility of horses, and it can be said that horses with larger keel values are runner forms that adapt to more open environments (Eisenmann, 1995; Koufos, 2016).

#### 5. Systematic Paleontology

Ordo : ARTIODACTYLA Owen, 1848

Family : BOVIDAE Gray, 1821

Genus : *Skoufotragus* Kostopoulos, 2009

Type species : *Skoufotragus schlosseri* (Andrée, 1926)

*Skoufotragus laticeps* (Andrée, 1926)

**Age:** early-middle Turolian (MN12) (late Miocene)

**Material:** Upper left M<sup>1-2</sup> (EUNHM-PV-8000) (Figure 5A)

**Description:** The specimen represents an upper left molar series of a semi-adult individual (length x width (mm): M1:20.3x19.2; M2:23.5x20.1). In both molars, entostyle is not observed with junction of the anterior and posterior medial lobes on the chewing surface. The protocon is angular and pointed, while the hypocon is larger and rounded. The basal pillar is not observed in the middle of the chewing surface of the tooth. Metastyle is the weakest and parastyle is the most developed style. M2 is larger than M1 and the proximal of the tooth is wider than its distal. Parastyle and mesostyle are equally developed.

**Correlation/Comparison:** ‘Protoryxoid’ bovids have been recorded in the SE Mediterranean region from the late-middle Miocene onwards (Gentry, 2000; Kostopoulos and Karakütük, 2015). During the late Miocene, especially in Anatolia, Greece and its environs, the group was widespread and many taxa co-exist in the same age. The taxonomy of the genera

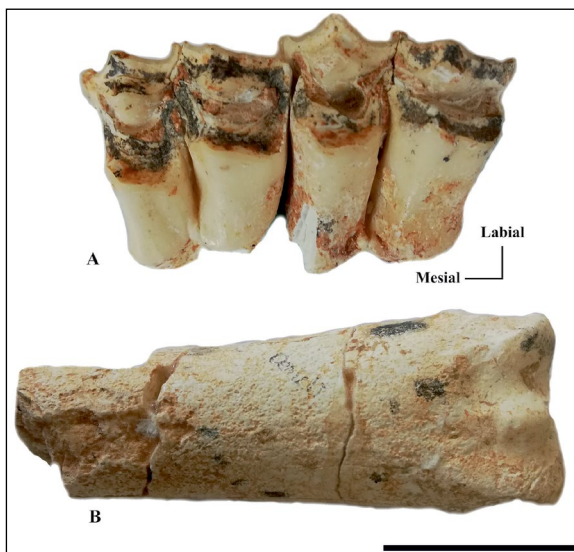


Figure 5- Fossil specimens of Denizli-Gölyeri fauna:

A. *Skoufotragus laticeps* upper left M<sup>1-2</sup> (EUNHM-PV-8000)

B. *Hipparion brachypus*: left distal MC3 (EUNHM-PV-8001) (Scale: 5a: 4 cm, 5b: 2 cm)

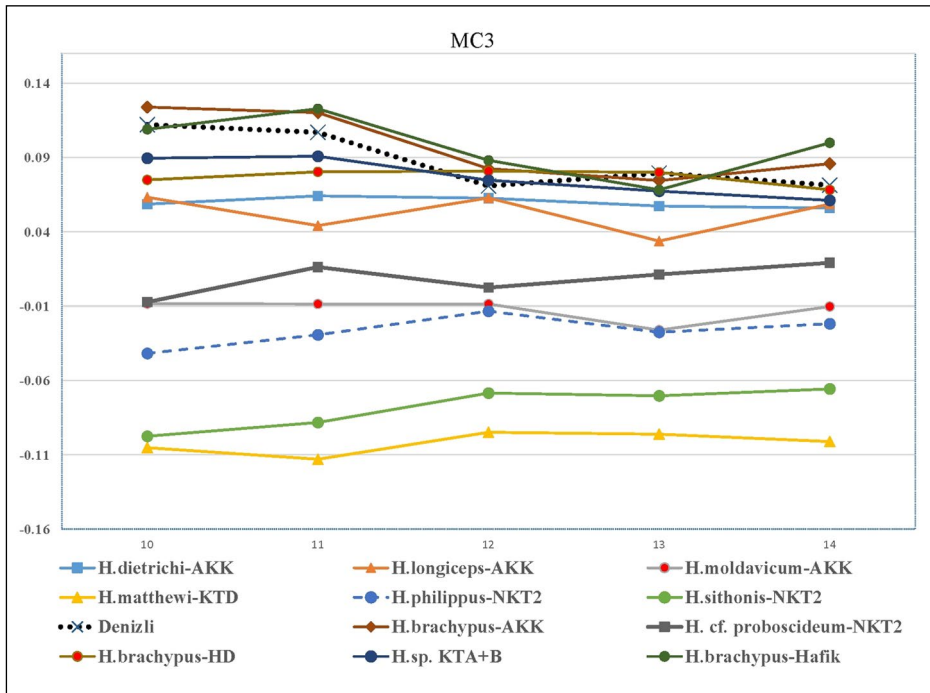


Figure 6- The logarithmic comparison diagram of the Denizli-Gölyeri fauna and various late Miocene *Hipparion* species presented by Hristova et al. (2003), Koufos and Kostopoulos (1994) and Koufos and Vlachou (2005) (Standard: *H. mediterraneum*, Pikermi-Greece, Koufos 1987). The M10-M14 measurements are compared with the findings by Hristova et al. (2003), Koufos and Kostopoulos (1994) and Koufos and Vlachou (2005).

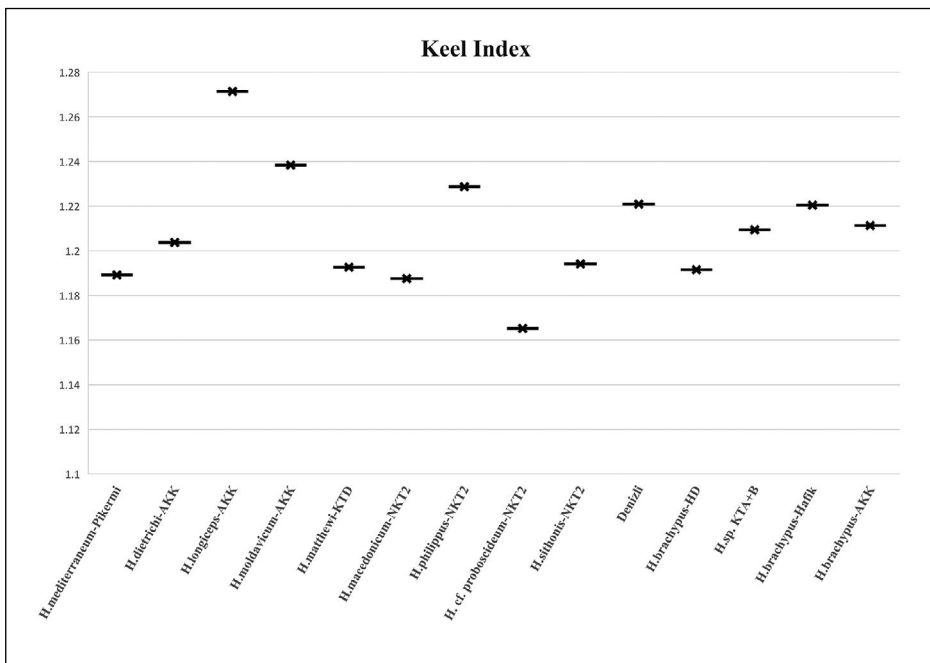


Figure 7- The MC3 keel index diagram including Denizli-Gölyeri fauna and various late Miocene *Hipparion* species presented by Hristova et al. (2003), Koufos and Kostopoulos (1994) and Koufos and Vlachou (2005).

involved was revised by Kostopoulos (2009), who suggested replacing the problematic *Pachytragus* by the new name *Skoufotragus*, keeping *Protoryx* only for *Protoryx carolinae* and *Pr. capricornis*. In Anatolia *Skoufotragus* is recorded in the localities of Kınık (*Skoufotragus* sp., late Miocene, MN12) (Köhler, 1987), Kemiklitepe-A (*S. laticeps*, late Miocene, MN11) (Bouvrain, 1994), Akkaşdağı (*S. schlosseri*, late Miocene) MN12) (Kostopoulos, 2009) and Şerefköy-2 (*S. cf. schlosseri*, late Miocene, MN12) (Kostopoulos and Karakütük, 2015). In addition, this genus has been recorded in Greek (Samos, Kryopigi) (Kostopoulos, 2009; Lazaridis et al., 2018) and Iranian (Maragheh) localities (Kostopoulos and Bernor, 2011).

The Gölyeri specimen is similar to *S. laticeps* morphologically (absence of entostylid, narrow protocon in M1, relatively high hypsodonty (M1:102, M2:110 mm height/width) and biometrically, and larger than *S. schlosseri* and *S. zemalisorum* samples from neighbouring localities (Kostopoulos, 2009; Kostopoulos and Bernor, 2011).

Ordo : PERISSODACTYLA Owen, 1848

Family : EQUIDAE Gray, 1821

Genus : *Hipparion* de Christol, 1832

*Hipparion brachypus* Hensel, 1862

**Age:** early-middle Turolian (MN12) (late Miocene)

**Material:** Left distal metacarpal 3 (MC3) (EUNHM-PV-8001) (Figure 5B) (M10: 43.57, M11: 41.15; M12: 31.5; M13: 25.8; M14: 28.53)

**Description:** The sample is a distal part of a large equid metacarpal. It is highly compatible with *H. brachypus* samples in size (Figure 6). In addition, the keel index (M12×100/M13) (Staesche and Sondaar, 1979; Şen et al., 1978, Eisenmann et al., 1988) value is compatible with *H. brachypus* values known from Anatolia and nearby regions (Akgün et al., 2000; Kaya et al., 2012) (Figure 7).

In Anatolia, large forms of late Miocene Equidae findings were grouped under *H. brachypus* species. To date, these findings are known in the localities of Kemiklitepe-A (late Miocene, MN11) (Koufos and Kostopoulos, 1994), Akkaşdağı (late Miocene,

MN12) (Koufos and Vlachou, 2005), Şerefköy-2 (late Miocene, MN12) (Kaya et al., 2012) and Sivas-Hafik (late Miocene, MN12) (Akgün et al., 2000) (Figure 8). The species is also present in the late Miocene localities of Greece (Samos, Pikermi) (Koufos and Vlachou, 2005; Vlachou and Koufos, 2009), Bulgaria (Hadjidimovo, Kalimanci) (Hristova et al., 2003) and Iran (Maragha) (Bernor et al., 2016). The relatively short and coarse metapodials suggest that this species was adapted to living in more confined paleo-environments (Koufos and Vlachou, 2005; Vlachou and Koufos, 2009).

## 6. Discussion and Conclusions

The mammalian fossils known in the Neogene succession of the Denizli Basin have shown that the region is an important migration route for many terrestrial taxa during the Neogene period. In this study, the Turolian stage is defined for Denizli Basin for the first time in the regional paleobiogeographic context. These fossils are located in the Güzelpınar member which is composed of alluvial fan deposits and defined during this study within the Kolankaya formation which is tectonically uplifted on the northeastern margin of the basin. The Gölyeri vertebrate fossil fauna is composed of taxa within the Anatolia and Greco-Iranian bioprovince extending from Greece to Iran throughout the late Miocene (MN11-12, early-middle Turolian). Similar taxa are known in the context of Turolian faunas from Asarlık, Mahmutgazi, Sazak, Elmalyurt (Pırnaz), Kemiklitepe and Karabeyli, Şerefköy, Özlüce and Karamusalar localities around Denizli Basin. Although the Gölyeri fauna contains a small number of specimens, it is of particular importance in terms of the stratigraphic correlation between the basins and the understanding of regional paleobiogeographic distributions as the first Turolian evidence in the Denizli Basin.

## Acknowledgement

This study was framed within the Paleoseismology Research Project of Turkey supported by General Directorate of Mineral Research and Exploration, Directorate of Geology Research Department with grant number of 2018-30.14.05 We would like to thank the project director Hasan Elmacı, S.M. was supported by Ege University, İzmir (research projects:



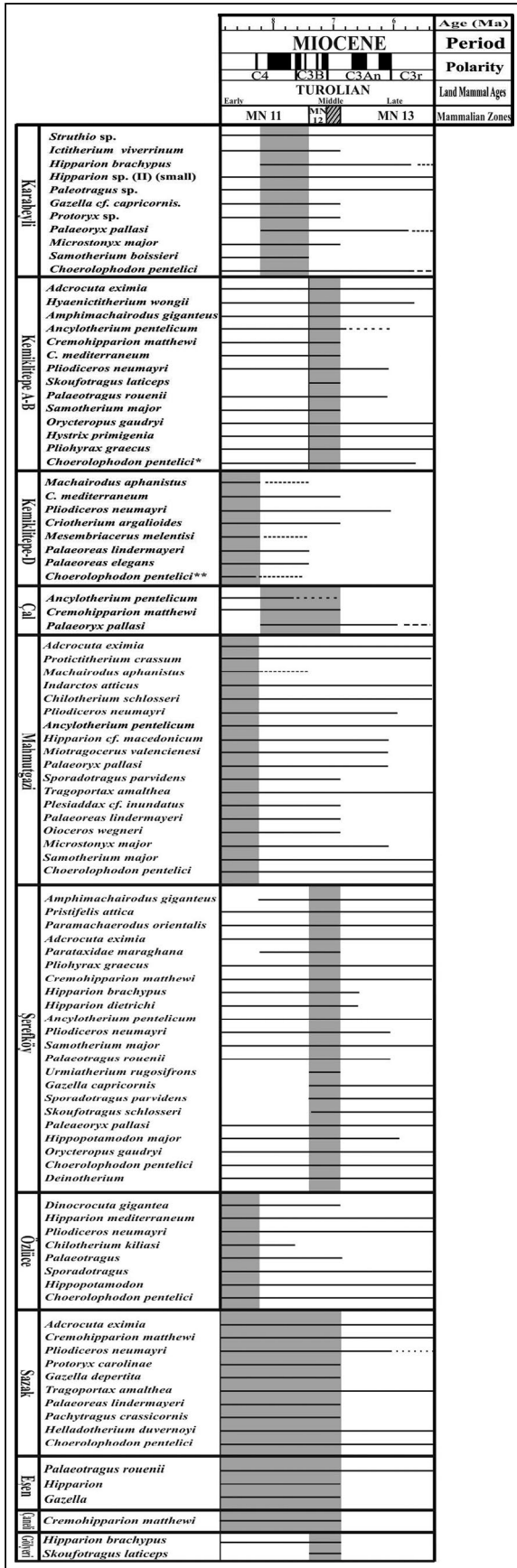


Figure 8- Biostratigraphic distribution and correlation of Turolian localities known in western Anatolia. Karabeyli (Seyitoğlu et al., 2009), Kemiklitepe (Şen et al., 1994; Saraç, 2003; Xafis et al., 2020), Çal-Asarlık (Alçiçek et al., 2012, Alçiçek and Alçiçek, 2014, Alçiçek et al., 2019), Mahmutgazi (Sickenberg et al., 1975; Pickford and Ertürk, 1979; Köhler, 1987; Saraç, 2003; Pickford, 2016; Geraads, 2017), Şerefköy (Saraç, 2003; Kaya et al., 2012; Mayda, 2014), Özlüce (Saraç, 2003; Alpagut, et al. 2014; Mayda, 2014), Sazak (Saraç 2003, Oruç, 2009, Tuna, 1999, Kaya, 1993), Eşen-Karamusalar (Alçiçek, 2007; Alçiçek et al., 2019), Çameli-Elmalıyurt (Saraç, 2003; Alçiçek et al., 2005, 2006, 2019), Sazak (Saraç, 2003; Tuna, 1999), Gölyeri (in this study).

TTM/001/2010, TM/001/2013, TTM/001/2014, TTM/001/2016, TTM/002/2016) during field and lab works. Constructive comments by T. Tanju Kaya (Ege Univ.) are appreciated. The manuscript was critically read and improved by Dimitris S. Kostopoulos (Aristotle University of Thessaloniki).

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